

Off-fault damage characterisation during and after experimental quasi-static and dynamic rupture in crustal rock from laboratory P -wave tomography and microstructures.

Franciscus M. Aben,¹ Nicolas Brantut¹, and Thomas M. Mitchell¹

Abstract. Elastic strain energy released during shear failure in rock is partially spent as fracture energy Γ to propagate the rupture further. Γ is dissipated within the rupture tip process zone, and includes energy dissipated as off-fault damage, Γ_{off} . Quantifying off-fault damage formed during rupture is crucial to understand its effect on rupture dynamics and slip-weakening processes behind the rupture tip, and its contribution to seismic radiation. Here, we quantify Γ_{off} and associated change in off-fault mechanical properties during and after quasi-static and dynamic rupture. We do so by performing dynamic and quasi-static shear failure experiments on intact Lanhélin granite under triaxial conditions. We quantify the change in elastic moduli around the fault from time-resolved 3D P -wave velocity tomography obtained during and after failure. We measure the off-fault microfracture damage after failure. From the tomography, we observe a localised maximum 25% drop in P -wave velocity around the shear failure interface for both quasi-static and dynamic failure. Microfracture density data reveals a damage zone width of around 10 mm after quasi-static failure, and 20 mm after dynamic failure. Microfracture densities obtained from P -wave velocity tomography models using an effective medium approach are in good agreement with the measured off-fault microfracture damage. Γ_{off} obtained from off-fault microfracture measurements is around 3 kJm² for quasi-static rupture, and 5.5 kJm² for dynamic rupture. We argue that rupture velocity determines damage zone width for slip up to a few mm, and that shear fracture energy Γ increases with increasing rupture velocity.

1. Introduction

During shear failure in rock, stored elastic strain energy is partly released as radiated energy E_r (i.e., seismic waves) and mostly dissipated on and around the fault interface as latent heat and new fracture surface area through a plethora of dissipative processes. Dissipated energy is typically partitioned into frictional work and breakdown work, where frictional work E_f is the work done to overcome the residual friction on the fault interface during sliding. Breakdown work W_b is a collective term of energies dissipated in addition to E_f , and primarily includes dissipative processes that reduce the strength of the fault interface towards the residual friction. This includes comminution, flash heating [Brantut and Viesca, 2017], and thermal pressurisation [Viesca and Garagash, 2015], but also includes energy dissipated towards propagating the rupture tip, and energy dissipated by deformation outside the principal slip zone (off-fault deformation). For earthquakes, E_r and W_b can be determined from seismological data [Tinti *et al.*, 2005; Kanamori and Rivera, 2006], where W_b varies from 10² to 10⁸ Jm⁻² as a function of total coseismic slip [Abercrombie and Rice, 2005; Viesca and Garagash, 2015]. As the strength evolution of the fault during failure cannot be determined directly from seismological data, a slip-weakening law is typically assumed to determine a slip-weakening distance δ_0 , at which the fault has reached its residual frictional strength. Seismological estimates for W_b do not discriminate between energy dissipated to propagate the rupture, Γ , and the remaining breakdown work ($W_b - \Gamma$). Γ is called the shear fracture

energy and is the energy dissipated within a process zone surrounding the rupture tip to overcome cohesion of the material and propagate the rupture by a unit area [Freund, 1990]. Γ is dissipated in a volume around the rupture tip, and may therefore include an off-fault component Γ_{off} in addition to the component of Γ dissipated to form the fault interface or principal slip zone. Measurements for material parameter Γ are of the order of 10⁴ Jm⁻² for initially intact crystalline low porosity rock under upper crustal conditions [Wong, 1982, 1986; Lockner *et al.*, 2001; Aben *et al.*, 2019], which may be considered an upper bound for pre-existing fault zones often comprised of damaged and altered rock. As Γ is dissipated earliest during shear failure [Barras *et al.*, 2020], its constituent dissipative processes may affect slip weakening processes in the wake of the rupture tip process zone – and may affect the remainder of W_b . We here aim to quantify the off-fault component of the fracture energy, Γ_{off} .

Off-fault deformation during shear failure, mainly fracturing and subsidiary slip, is created by transient off-fault stresses near the rupture tip [Andrews, 1976; Poliakov *et al.*, 2002; Rice *et al.*, 2005] and by increasingly larger off-fault stresses arising from progressive slip along rough faults [Chester and Chester, 2000; Dieterich and Smith, 2009]. Energy dissipated by off-fault deformation in the rupture tip process zone Γ_{off} is one of Γ 's constituent energy sinks. During shear failure, off-fault deformation caused directly by the stress concentration around the rupture tip as part of Γ_{off} precedes most of the off-fault deformation from slip on a rough fault, since the amount of slip within the rupture tip process zone is negligible. Off-fault deformation, and particularly off-fault fracturing, changes the mechanical and hydraulic properties of fault damage zone rock, and thus the constituent dissipative processes of Γ_{off} affect fault damage zone properties at an early stage during shear failure [Aben *et al.*, 2020]. This can have a feedback on rupture, slip, and ground motion; rupture simulations show that reduced mechanical properties in the fault damage zone affect fault slip [Cappa *et al.*, 2014] and slip velocity [Andrews, 1976, 2005; Dunham *et al.*, 2011]. Due to fracturing near the rupture tip the pore volume increases and causes, under partially undrained conditions, a local pore fluid pressure drop and an increase in effective pressure on the fault [Brantut, 2020]. This can stabilise dynamic rupture [Martin, 1980] and

¹Department of Earth Sciences, University College London, London, UK

This version is a home made pdf generated from the original L^AT_EX files, and the formatting and editing differs from the version published by AGU. Minor differences with the official version may subsist (British spelling, figure numbering).

slip [Segall and Rice, 1995; Segall et al., 2010]. Changes in hydraulic properties from off-fault fracture damage close to the fault interface have an effect on slip-weakening mechanisms that act in the wake of the rupture tip, such as thermal pressurisation [Brantut and Mitchell, 2018]. The dynamic reduction of elastic moduli in the fault damage zone causes high frequency content in the radiated ground motion [Thomas et al., 2017], and can be a substantial additional source of seismic radiation [Ben-Zion and Ampuero, 2009]. It is therefore crucial to 1): Quantify Γ_{off} , and 2): Quantify the changes it imposes on off-fault mechanical properties.

A measurement of total off-fault fracture surface area created in the rupture tip process zone gives an estimate for the cumulative fracture surface energy necessary to create them. This gives a lower bound for Γ_{off} , as energy dissipated as latent heat during off-fault fracturing (i.e., slip on the fractures) remains unknown. Along strike-slip faults, this approach has yielded estimates for the total off-fault dissipated energy [Chester et al., 2005; Rockwell et al., 2009]. However, fractures observed in exhumed fault damage zones originate from either rupture tip stress concentrations, stresses generated by slip on a rough fault during shear failure, or quasi-static stresses [Mitchell and Faulkner, 2009], and were healed and overprinted by numerous shear failure events. This complicates quantification of Γ_{off} from the geological record. Off-fault fracture damage induced by shear failure under controlled conditions in the laboratory circumvents some of these complications, allowing for a microstructural description [Wawersik and Brace, 1971; Reches and Lockner, 1994] and quantification of fracture damage zones [Moore and Lockner, 1995; Zang et al., 2000] associated to a single failure event. Moore and Lockner [1995] estimated the cumulated surface energy in the fracture damage zone around a ‘frozen’ quasi-static rupture front in granite, where slip on the fault was negligible, yielding a lower bound for Γ_{off} . A dynamically propagating rupture tip is expected to create a larger area of fracture damage, as the stress field around a propagating rupture tip is distorted with increasing rupture velocity [Poliakov et al., 2002], and we therefore expect Γ_{off} to increase as well.

Γ_{off} can also be obtained from the change in stored elastic strain energy in the rupture tip process zone, with the underlying assumption that the change in elastic compliance is caused by off-fault fracturing. A reduction in elastic compliance is measured as a drop in seismic wave speeds, making them an attractive and cost-efficient proxy for large scale monitoring of fracture damage structures in fault zones [Mooney and Ginzburg, 1986; Rempe et al., 2013; Hillers et al., 2016; Qiu et al., 2017]. To date, high resolution geophysical measurements of wave speeds from dense arrays [Ben-Zion et al., 2015] have given static snapshots of the fault damage zone structure, but not the coseismic velocity drop necessary to obtain the total coseismic off-fault dissipated energy, let alone Γ_{off} . Laboratory-scale seismic tomography of the P -wave velocity structure [Brantut, 2018] obtained from ultrasonic data measured during quasi-static shear failure experiments does give the change in effective elastic moduli during rupture needed to calculate Γ_{off} [Aben et al., 2019], yielding a similar value for Γ_{off} to that calculated from fracture surface area by Moore and Lockner [1995]. There are, to our knowledge, no measurements of Γ_{off} for dynamic shear ruptures yet, either from microstructures or from a change in elastic moduli.

The changes in off-fault mechanical properties induced by shear rupture cannot be assessed directly from the scalar quantity Γ_{off} , but the two approaches outlined above to estimate Γ_{off} also provide the changes in elastic moduli and the microfracture density. These two physical properties can be reconciled using effective-medium theory models for cracked solids [e.g., Guéguen and Kachanov, 2011], which are an important tool for obtaining information on physical and hydraulic properties such as fracture density [Sayers and Kachanov, 1995], porosity, and permeability [Gavrilenko and Guéguen, 1989]. These physical parameters are key in studying the feedback between rupture and slip. Effective-medium approaches have been tested in the laboratory on deformed samples, where effective elastic moduli were measured by active ultrasonic surveys [e.g., Schubnel et al., 2003]. The path-averaged wave velocities

obtained from these surveys are representative for fracture damage only when fractures are homogeneously spread throughout the sample. In laboratory shear failure experiments, fracture damage is localised around the fault interface and so path-averaged velocities cannot be used. Instead, recent advances in syn-deformation laboratory tomography techniques [Brantut, 2018; Stanchits et al., 2003] can be employed for the use of effective-medium models, so that changes in physical properties can be quantified in situ.

Here, we assess Γ_{off} for dynamic and quasi-static rupture in granite following the two approaches outlined above. To do so, we perform three types of shear failure experiments in the laboratory: Shear failure by quasi-static rupture, by dynamic rupture, and by partly quasi-static and partly dynamic rupture (from here on referred to as ‘mixed rupture’). We quantify the change in mechanical properties around the fault caused by shear failure from time-resolved 3D P -wave velocity tomography models. These were obtained during and after quasi-static rupture and after dynamic rupture and mixed rupture. We also quantify the off-fault microfracture damage after dynamic and quasi-static shear failure from microstructural observations. An effective-medium approach is used to obtain microfracture densities from the 3D P -wave velocity models, which are compared with the measured microfracture densities. We then determine a damage zone width for the quasi-statically and dynamically failed samples. These estimates for damage zone width are compared to the expected damage zone width from the stress field around a propagating rupture tip [Poliakov et al., 2002] and from the off-fault stresses induced by slip along a rough fault [Chester and Chester, 2000]. We then obtain Γ_{off} from measuring the cumulative off-fault fracture surface energy within the damage zone. These measurements are complementary to Γ_{off} derived from changes in effective elastic moduli by Aben et al. [2019] for a quasi-static rupture. Last, we discuss the implications of our results to the energetics of earthquake rupture.

2. Materials and methods

2.1. Experiments

Three different types of failure experiments were performed on intact 100 mm by 40 mm diameter Lanhélin granite cylinders (from Brittany, France) at 100 MPa confining pressure (Table 1): Failure by dynamic rupture, failure by quasi-static rupture, and failure by part quasi-static rupture and part dynamic rupture named mixed rupture. The experiments were performed at nominally dry conditions in a conventional oil-medium triaxial loading apparatus at University College London [Eccles et al., 2005]. Axial load was measured by an external load cell corrected for friction at the piston seal. Axial shortening was measured by a pair of Linear Variable Differential Transducers (LVDTs) outside the confining pressure vessel, corrected for the elastic shortening of the piston.

The samples were equipped with two pairs of axial-radial strain gauges. The samples were placed in a rubber jacket equipped with 16 piezoelectric P -wave (V_p) transducers. Ultrasonic signals were amplified to 40 dB before being recorded by a digital oscilloscope (50 MHz sampling frequency). All signals consisted of 4096 data points, equivalent to an 82 μs time interval. Active ultrasonic velocity surveys were performed every 5 minutes, where all 16 piezoelectric transducers were sequentially used as a source, while the other transducers recorded the resulting waveforms. 1 MHz pulses were produced by exciting the source transducer with a 250 V signal. The signal-to-noise ratio was improved by stacking the recorded waveforms from six of these pulses per transducer. Between surveys, acoustic emissions (AE) were recorded on 16 channels, provided that the AE signal amplitude was above 250 mV on at least two channels within a 50 μs time interval. The digital oscilloscope stored up to four sets of AE waveforms per second.

Dynamic rupture was achieved by setting a constant shortening rate equivalent to an axial strain rate of 10^{-5} s^{-1} until dynamic shear failure. Quasi-static rupture was achieved by suppressing dynamic rupture via monitoring the AE rate, following the approach of Lockner et al. [2001]. When the acoustic emission rate showed

Table 1: Sample table with experiment type, number of time intervals, number of AE events used for tomographic inversion, tomographic inversion parameters (covariances and correlation length), and number of SEM images used for microstructural analysis. [†]From *Aben et al.* [2019]. [‡]Accumulated slip by reloading after failure.

sample number	LG1	LN4	LN5 [†]	LN7	LN8
type of experiment	dynamic [‡]	mixed	quasi-static	dynamic	mixed [‡]
nr. of time intervals	8	-	38	-	22
number of AE events	2215	-	11134	-	9844
survey arrival time [μ s]	1	-	1	-	1
anisotropy parameter [-]	0.01	-	0.01	-	0.01
<i>a priori</i> velocity model [log(m/s)]	0.02	-	0.02	-	0.01
AE arrival time [μ s]	2	-	2	-	2
AE source location [mm]	2	-	2	-	2
AE origin time [μ s]	2	-	2	-	2
correlation length [mm]	25	-	25	-	25
microstructural analysis	-	-	123 images	134 images	-
slip δ [mm]	2.88	1.93	0.83	3.22	2.44

a marked increase – a precursor to dynamic rupture – the axial load on the sample was decreased by reversing the displacement direction of the piston. For mixed rupture experiments, the rupture was controlled for about half the stress drop between the sample’s peak stress and its residual frictional strength. The rupture was allowed to propagate dynamically for the remainder of the stress drop. After failure, one sample failed by dynamic rupture and one sample failed by mixed rupture were reloaded up to their residual frictional strength (Table 1), which resulted in some additional stable sliding along the fault.

Poisson’s ratio of the intact rock ν_0 was determined from the ratio of the axial and radial strain during axial loading in the elastic regime. The intact Young’s modulus E was derived from the differential stress versus axial displacement curves measured during axial loading in the elastic regime.

2.2. Analysis of ultrasonic data and P -wave tomography

The *FaATSO* code by *Brantut* [2018] was used for tomographic inversion of the active ultrasonic surveys and AE arrival times. Prior to tomographic inversion, the ultrasonic waveforms recorded during the experiments were processed. Time of flight for all sensor combinations were picked for the first active ultrasonic survey of the experiment, and arrival times for subsequent surveys were extracted using an automated cross-correlation technique [e.g., *Brantut et al.*, 2014] with a precision of about 0.05 μ s. From these, path-averaged velocities were calculated between sensor pairs. These ray paths are oriented at 90° (i.e., horizontal), 58°, 39°, and 28° angles to the loading axis of the sample. AE arrival times and source locations were obtained in three steps: 1) The first arrivals of the AE waveforms were automatically picked, and AE source locations were calculated using their arrivals in conjunction with a transverse isotropic velocity model based on the most recent ultrasonic survey. 2) The AE events were subjected to a quality test, where AEs with a source location error above 5 mm were discarded. 3) The automatically picked arrival times of the remaining AEs were subjected to an interactive visual check – arrival times were improved or removed when the difference between the automatically picked arrival time and the theoretical arrival time for the calculated source location was too large. 4) The AE source locations were recalculated based on the inspected arrival time dataset and the same source location error criterium was applied.

The *FaATSO* code treats the arrival times of the ultrasonic surveys and the AE arrival times as the observed data. The model parameters are the AE source locations and origin times, and the horizontal P -wave velocity and anisotropy in voxels of $5 \times 5 \times 5$ mm that cover the sample volume. The algorithm allows for vertical transverse isotropy for each voxel (i.e., the vertical velocity is independent from the horizontal velocity). V_P anisotropy is expressed as the ratio $(V_P^v - V_P^h)/V_P^h$, where V_P^h and V_P^v are the horizontal and vertical P -wave velocities, respectively. To make predictions of the observed data based on the model parameters, a 3D anisotropic

ray tracer is used (i.e., Eikonal solver) [*Brantut*, 2018]. The inverse problem is solved using a quasi-Newton inversion algorithm [*Tarantola*, 2005], and is constrained by a set of standard deviations that describe Gaussian variances on the observed data (Table 1). The variance on the model parameters (AE source locations, velocity, and anisotropy) are also Gaussian, expressed by standard deviations (Table 1). For the velocity and anisotropy, there is a covariance between voxels that is a function of the variance and a correlation length [*Brantut*, 2018]. Through the covariance for velocity and anisotropy, the correlation length smooths heterogeneities in the inversion results.

The observed data was divided in a number of time intervals (Table 1) with varying duration, each containing roughly 300 AEs, for which we performed the inversion. The AE source locations were used as an *a priori* model parameter. For the remaining *a priori* model parameters, V_P^h and anisotropy, we used two structures: 1) A homogenous vertical transverse isotropic (VTI) *a priori* velocity structure derived from the most recent ultrasonic survey in each time interval, and 2): An inherited *a priori* velocity structure from the inversion results of the preceding time interval, except for the first time interval where a homogenous VTI *a priori* structure was used. The quality of the inversion results was tested by comparing both sets of inversion results (see Text S1).

A clear tomographic image during dynamic rupture could not be achieved by inversion of pre- and syn-rupture AE events, because the number of recorded syn-rupture events is too low due to the limited recording capacities of the acquisition system, and pre-rupture events occur at a stage where deformation is not yet localised. We therefore use AE events recorded during reloading of a dynamically failed sample and a sample failed by mixed rupture.

2.3. Microstructural analysis

Polished thin sections oriented perpendicular to the main fault interface were cut from epoxied post-mortem samples that failed by dynamic rupture and by quasi-static rupture. The thin sections were studied by optical microscopy and by scanning electron microscope (SEM), from the latter we obtained back-scatter electron (BSE) grayscale images along three transects through the centre of the sample (Figure 1a, b; Table 1). The images were taken at a $100\times$ magnification and cover a 1.0 mm^2 area. The pixel dimension is 0.5 by $0.5 \mu\text{m}$.

We obtained the traces of microfractures as follows: Microfractures are revealed as low grayscale value features in the SEM pictures, because they are empty or filled with low density epoxy. The microfractures may be traced by hand, but given the large number of SEM images, we elected to use a semi-automated image analysis technique. Both methods are prone to user errors, but the errors from semi-automated image analysis are more consistent in all images so that analysis within the dataset itself is more reliable. The microfractures can be isolated by using a grayscale threshold, but this approach will isolate pores in addition to open

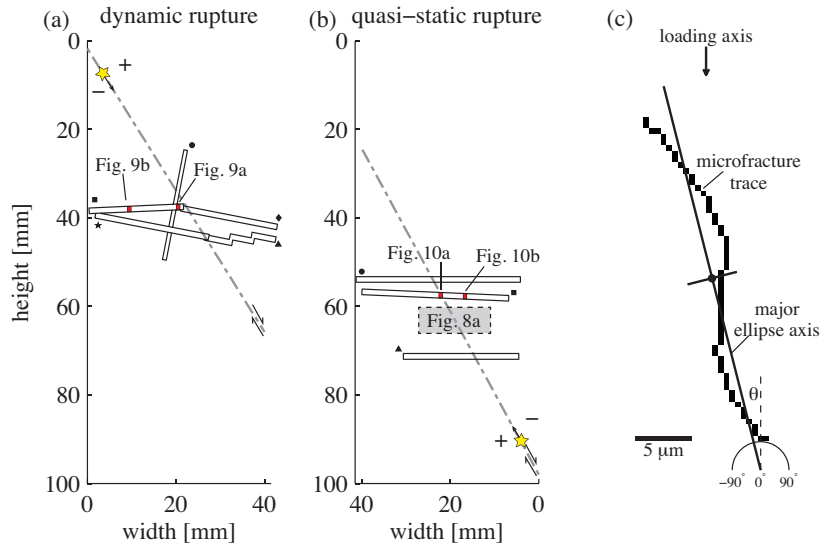


Figure 1: Sketch of fault-perpendicular slice through the centre of a sample showing locations of the transects along which SEM images have been obtained (to scale), for (a): Dynamic rupture, sample LN7, and (b): Quasi-static rupture, sample LN5. Each transect is assigned a symbol corresponding to the fracture density data in Figures 9d and 10d. The trace of the main fault plane is shown by the dashed gray line. The red windows indicate the locations of individual SEM images shown in Figures 9a, b, and 10a, b. The compressive and tensile lobes in the rupture tip process zone are marked by + and - signs, the star and arrow on the fault marks the approximate location of rupture nucleation and propagation direction. Location of the optical microscopy image in Figure 8a indicated in (b). (c): Trace of a fracture segment, and the minor and major axis of an ellipse fitted around the segment. The fracture segment length used to calculate ρ_{frac} is given by the number of its constituent pixels. The angle θ between the major ellipse axis and the loading axis gives the fracture orientation.

fractures, and will exclude pixels of low aperture fractures because they partly overlap with higher density wall rock, which increases the absolute grayscale value. Fracture recognition from sharp grayscale contrasts (i.e., edge detection) is more sensitive to low aperture fractures, but will also recognise pores and sharp grain boundaries between different minerals. Here, we isolate microfractures based on fracture aperture, so that larger aperture pores can be excluded. To do so, we use the median filter technique used by Griffiths *et al.* [2017], and incorporate their approach in the newly developed fracture tracing code *Giles* (fracture tracing by median filter, skeletonisation, and targeted closure, freely available on https://github.com/FransMossel/Giles_fracturetracing.git). The median filter obtains a median grayscale value for a predefined window of pixels around a target pixel, and assigns this median value to the target pixel. The entire image is subjected to this action. If the predefined window is larger than the fracture aperture and smaller than the aperture of pores, it ascribes a median grayscale value to a pixel in the fracture that is much higher than the original value, but pixels that represent pores or grains do not significantly change [Griffiths *et al.*, 2017]. The difference between the original grayscale values and the median filtered values is thus much higher in microfractures than in surrounding grains and pores. The image of this difference is therefore binarised. Small gaps between fracture traces in the binarised image are closed with a dilation-erosion action. The binary image is skeletonised, reducing the width of the trace to a single pixel, followed by targeted closure of gaps between traces that have the same orientation. Small residual branches on the fracture traces are an artefact of the skeletonisation process, and are removed by a pruning algorithm similar to that used by Griffiths *et al.* [2017]. A visual check and, when necessary, adjustment of the user-defined parameters, is imperative to ensure reasonable results from the fracture tracing code. See Text S1 for more details on the image analysis steps.

The end result of image processing using *Giles* is a binary image with microfracture traces of single pixel width. Fractures below 3 to 9 μm in length (depending on the size of the median filter window) were not traced. Since we are primarily interested in off-fault

damage, we manually removed fracture traces in gouge-filled zones and zones of cataclasite. We do not define individual fractures, because this requires manual unravelling of the microfracture network that would give arbitrary results for a well-connected fracture network where a clear fracture hierarchy is missing. Instead, we analyse fracture segments, which are defined as pixels connected to only two neighbours. Fracture segments are separated by fracture intersections, which are pixels with three or more neighbours.

We obtained 2D fracture orientations for each fracture segment by fitting an ellipse around a segment and measuring the angle θ between the major axis of the ellipse and the sample axis (Figure 1c). The absolute cumulative fracture length in an image is given by the number of pixels used for the fracture traces. Off-fault fracture density ρ^{frac} (in mm/mm^2) was obtained for each image by dividing the total fracture length in an image with the surface area of that image. The SEM image transects span both sides of the fault zone, which experienced different transient stresses in the rupture tip process zone. The transient off-fault stresses are tensile on the side of the fault where the direction of slip is opposite to the rupture propagation direction, and compressive on the other side of the fault. Based on the migration of AE source locations over time, which indicates the rupture propagation direction, we identified the tensile and compressive sides of the fault (Figure 1a, b).

3. Results

3.1. Experiments

The samples reached a peak differential stress of 660 to 700 MPa, followed by the onset of fault localisation and rupture propagation (Figure 2a). Frictional sliding – and thus the completion of rupture – commenced between 360 and 350 MPa, based on the flattening of the stress-displacement curve and the spread of the AE source mechanisms across the entire slip surface during quasi-static rupture. The post-failure residual strength is around

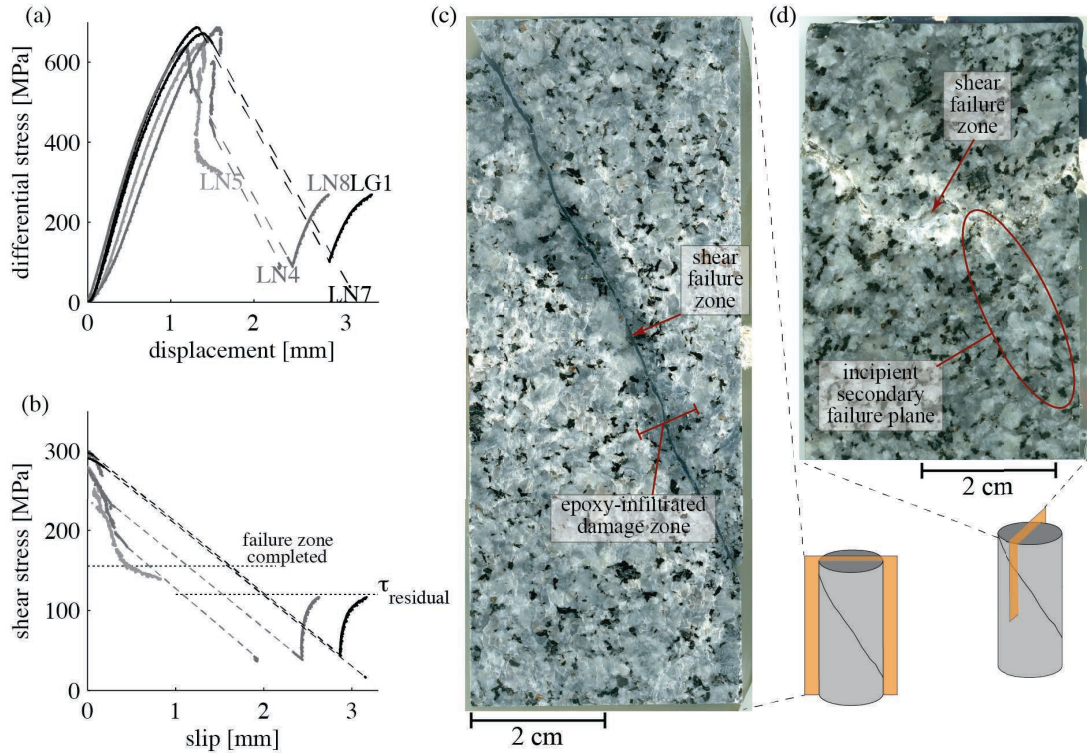


Figure 2: **(a)**: Stress-displacement curves for Lanhélin granite samples subjected to dynamic rupture (LN7 and LG1, black), to quasi-static rupture (LN5, light gray), and to mixed rupture (LN4 and LN8, dark gray). The displacement and stress drop caused by dynamic failure are shown as dashed intervals. **(b)**: Shear stress versus slip curves for all failure experiments. Curves follow the same colour coding as in (a). The shear stress at which the failure zone was completely formed is indicated, and $\tau_{residual}$ gives the residual shear stress to which the curves converge. **(c)**: Polished section of dynamically failed sample LN4, oriented perpendicular to the failure zone. Note that epoxy has penetrated the failure zone and part of the damage zone (darkened area), but has not penetrated the intact rock near the edge of the sample. **(d)**: Polished section of mixed ruptured sample LN8, oriented parallel to the main failure zone. A perpendicular incipient failure plane is visible. Surface has been epoxied prior to polishing, so that the damage zone is less apparent compared to (c).

300 MPa, as shown by the converging stress-displacement curves of the quasi-static, dynamic, and mixed experiments, of which the latter two approached the residual frictional strength from a lower differential stress by reloading of the sample.

Visual inspection of the samples after the deformation experiment revealed a single shear failure zone (Figure 2c), except for sample LN8 that includes an incipient secondary fault plane without noticeable displacement in addition to the through-going shear failure zone (Figure 2d). All through-going failure zones are oriented approximately at 30° relative to the compression axis. Using this fault angle, we resolved the average shear stress on the fault plane from the differential stress and confining pressure (Figure 2b). The rupture fully traversed the sample and completed the failure zone at about 155 MPa shear stress (Figure 2b), as measured from the quasi-static rupture, and the residual frictional strength $\tau_{residual}$ is around 120 MPa, as shown by the converging stress-strain curves for quasi-static, dynamic, and mixed ruptures (Figure 2b). The slip on the fault δ , calculated from the axial displacement data corrected for machine stiffness and for the stiffness of the intact rock, was 0.83 mm at the end of the quasi-static rupture experiment, 2.88 to 3.22 mm after dynamic failure, and 1.93 to 2.44 mm after dynamic failure in mixed rupture experiments (Table 1). Additional slip of 0.19 mm and 0.29 mm was accumulated by reloading samples LG1 and LN8, respectively, after dynamic failure.

A Young's modulus $E = 88$ GPa was measured for the intact rock during axial loading above 100 MPa and below about 400 MPa

differential stress, and averaged over all experiments. Averaged over all experiments, a Poisson's ratio $\nu_0 = 0.20$ was estimated for intact rock.

3.2. Ultrasonic velocity surveys

Path-averaged ultrasonic P -wave velocities were routinely calculated from the time of flight between two sensors, assuming a straight ray path (i.e., shortest distance) between the sensors. We present the P -wave velocity change during deformation with respect to the initial P -wave velocity at hydrostatic conditions along 5 straight ray paths at key orientations with respect to the fault plane during a quasi-static failure experiment (Figure 3a) and a dynamic failure experiment (Figure 3b). In both samples, ray path A is perpendicular to the loading axis and located well outside the eventual failure zone. Ray path B is oriented at 39° to the loading axis, and nearly its entire length is located within the failure zone. Ray paths C and D, both at a 58° to the loading direction, intersect with the two extremities of the fault zone and run sub-parallel to it. Ray path E, oriented perpendicular to the loading direction, intersects the fault zone in the centre of the sample.

Before the onset of quasi-static rupture, path-averaged P -wave velocities along all 5 ray paths increase slightly by 30 m/s up to 6.2 km/s from 0 to 400 MPa differential stress (Figure 3a), followed by a strong decrease as peak stress is approached. At the peak differential stress, V_p along ray path A shows the smallest velocity reduction of about 13% down to 5.3 km/s. During quasi-static rupture, when differential stress drops from peak stress to about

350 MPa, V_p along ray path A recovers by 6%, and remains stable during sliding between 350-300 MPa differential stress. Ray path B reveals a drop in P -wave velocity of about 13% at the peak differential stress. Between the peak stress and 600 MPa differential stress, V_p along ray path B decreases by an additional 4%, and remains stable at a total reduction of 17% for the remainder of the stress drop. After rupture completion at 350 MPa differential stress and the onset of sliding, the P -wave velocity along ray path B recov-

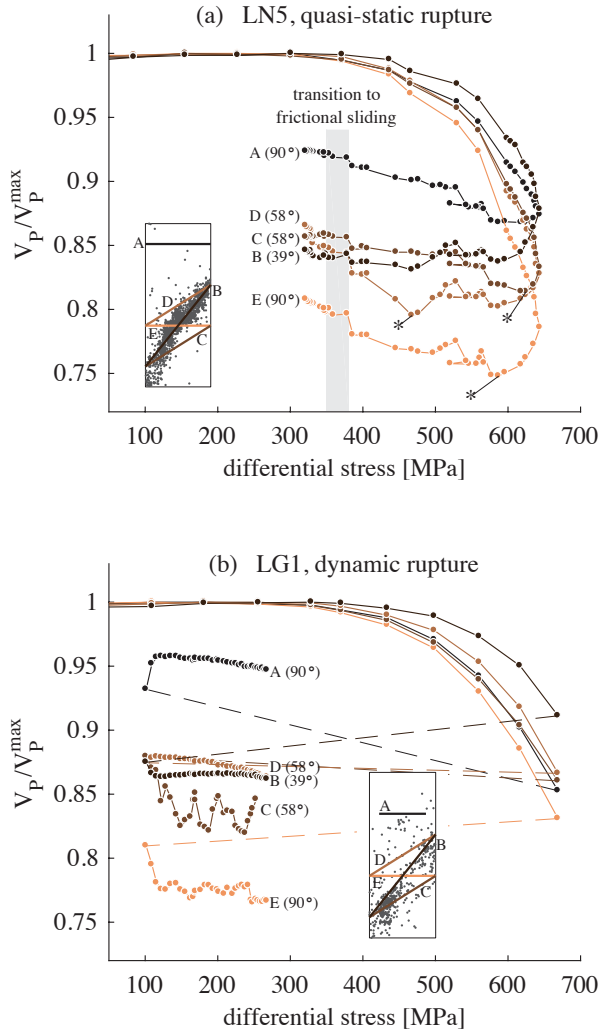


Figure 3: **(a)**: Normalised path-averaged P -wave velocity measured during a quasi-static failure experiment (sample LN5) versus differential stress. P -wave velocities were obtained from the first P -wave arrival of active ultrasonic surveys. The shaded area indicates the transition to frictional sliding, and the asterisks highlight the lowest velocities of three ray paths (see main text). The curves are coloured similar to their locations shown in the cross-section through the centre of the sample (inset) where the fault plane, delineated by AE source locations within 2 mm of the cross-section, intersects the cross-section and ray paths at a 45° angle. **(b)**: Normalised P -wave velocity measured before and after a dynamic failure experiment (sample LG1) versus differential stress. The dynamic stress drop during failure is dashed. The curves are coloured similar to their locations shown in the cross-section through the centre of the sample (inset) where the fault plane, delineated by AE source locations within 3 mm of the cross-section, intersects the cross-section and ray paths at a 40° angle.

ers by 1-2%. At the peak stress, P -wave velocity along ray path C and D dropped by 17%. V_p continues to decrease, along C down to 21% at 625 MPa differential stress, and along D down to 19% at 470 MPa (Figure 3a, asterisks). At the end of the experiment, after frictional sliding, the overall velocity drop along ray paths C and D is 14% and 15% respectively, which is a velocity recovery of 7% and 4% with respect to the minimum observed V_p . The velocity drop along ray path E was 22% at the peak stress. Along this ray path, we observe the strongest reduction in P -wave velocity of about 26% down to 4.6 km/s at a differential stress of 590 MPa during quasi-static failure (Figure 3a, asterisks). As failure progresses and differential stress drops further, the velocity recovers so that a 19% reduction in V_p is measured at the end of the experiment.

Path-averaged P -wave velocities before dynamic failure (Figure 3b) are similar to those before quasi-static rupture. Ultrasonic surveys could not be obtained during dynamic rupture, but were obtained during reloading of the sample after failure. P -wave velocity outside the fault zone along ray path A was reduced by 15% down to 5.3 km/s prior to dynamic failure from an initial velocity of 6.2 km/s. The dynamic stress drop during failure caused an increase in V_p of 8 to 11%, followed by a small decrease during reloading down to 5.9 km/s at 270 MPa differential stress (Figure 3b). P -wave velocity along ray path B drops from 5.7 km/s (9% drop) at peak stress to 5.4 km/s (13% drop) after dynamic failure (Figure 3b). During reloading, the P -wave velocity does not change along wave path B. Pre-failure P -wave velocities along ray paths C and D drop by 13–14% along both wave paths, and recover by 1% up to 5.4–5.5 km/s after the dynamic stress drop (Figure 3b). Within the fault zone along ray path E, the P -wave velocity drops during the dynamic stress drop from 5.1 km/s down to 4.8 km/s (22% reduction, Figure 3b). V_p decreases slightly more during reloading of the sample (down to a 23% reduction).

Path-averaged P -wave velocity changes measured during quasi-static rupture and before and after dynamic rupture are of similar magnitude and show a wide variation in velocity reductions within a single sample, with V_p reduced by 5 to 24% at the end of the experiment relative to the intact rock. These variations indicate strong localisation of damage, and the difference between horizontal V_p (measured perpendicular to the loading axis) and V_p measured at an angle indicate damage-induced anisotropy. Overall, P -wave velocity tends to increase with decreasing differential stress, except for the ray paths located entirely within the fault zone (ray paths B in Figure 3a and b). The above analysis assuming straight ray paths reveals very precise changes in path-averaged V_p thanks to the cross-correlation technique used to extract arrival times. However, changes in path-averaged V_p do not reveal where along the ray path the V_p has changed. Therefore, we perform a 3D tomographic inversion, which will lack the precision of the path-averaged velocity changes, but will reveal the location of greatest change in V_p .

3.3. P -wave tomography

We first present changes in the horizontal P -wave velocity structure introduced by quasi-static, dynamic, and mixed rupture. The results were obtained using an inherited *a priori* velocity model (see section 2.2 and Text S2). We then present the P -wave anisotropy inversion results. We detail the difference between the inherited and the homogeneous VTI *a priori* velocity models and the effect of different standard deviations on the model parameters in Supplementary information Text S2.

Dynamic rupture propagation: The horizontal P -wave velocity before localised dynamic failure drops throughout the entire sample, from 6.2 km/s down to 5.4 km/s (Figure 4a). The V_p structure obtained immediately after dynamic failure shows a strong localised low velocity zone (V_p drops by 22% down to 4.8 km/s) around the fault zone (Figure 4b). The localised zone of low V_p decreases in width from about 35 mm to 20 mm as the sample is reloaded. V_p recovers throughout the sample during reloading, most notably in the low velocity zone where the minimum P -wave velocity increases by 600 m/s to 5.4 km/s (Figure 4c, d).

Quasi-static rupture propagation: Before the onset of quasi-static rupture, horizontal V_p decreases from around 6 km/s down to 5 km/s (Figure 5a). The rupture nucleates near the bottom of the sample and propagates upwards (delineated by the AE source locations), during which a low P -wave velocity zone forms around

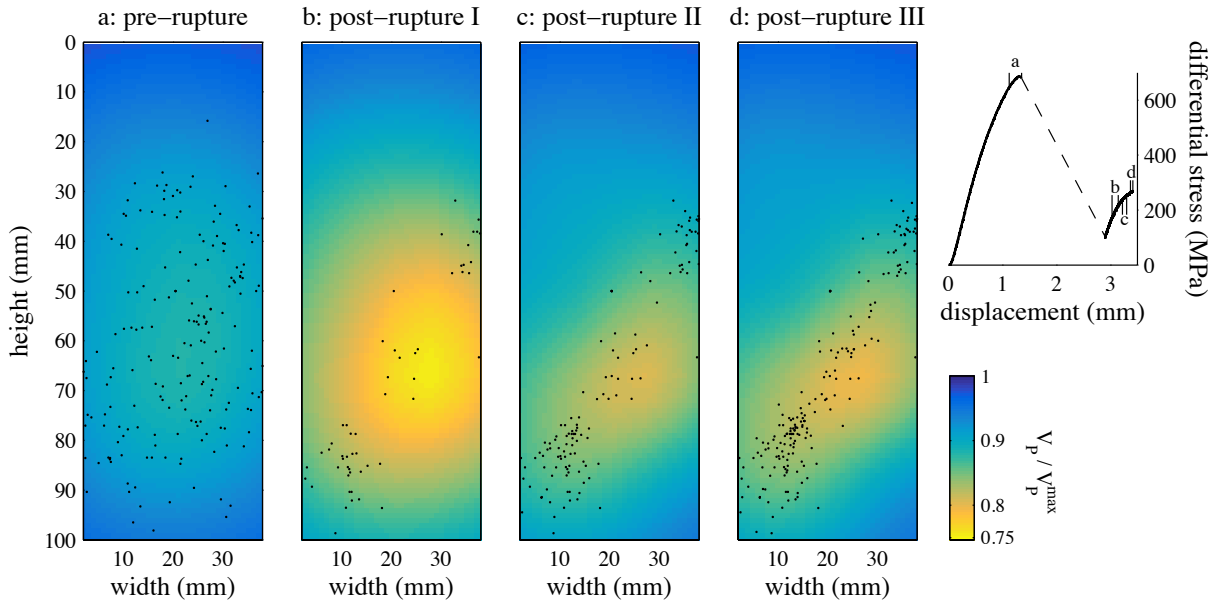


Figure 4: Tomographic sections of the horizontal V_p normalised to the initial velocity through the centre of sample LG1 (dynamic rupture), perpendicular to the fault. The four slices show time intervals **(a)** just prior to the localisation of deformation, and **(b)**, **(c)**, and **(d)** during three stages of reloading after rupture and slip. The corresponding parts of the stress-displacement curve are indicated on the right. AE source locations up to the time interval shown are projected onto the slice. The AE source locations are within 5 mm distance perpendicular to the slide. In this figure, and Figures 5 and 6, the seismic velocities are smoothed to a 1 mm resolution and an inherited *a priori* model was used for the inversions of ultrasonic data.

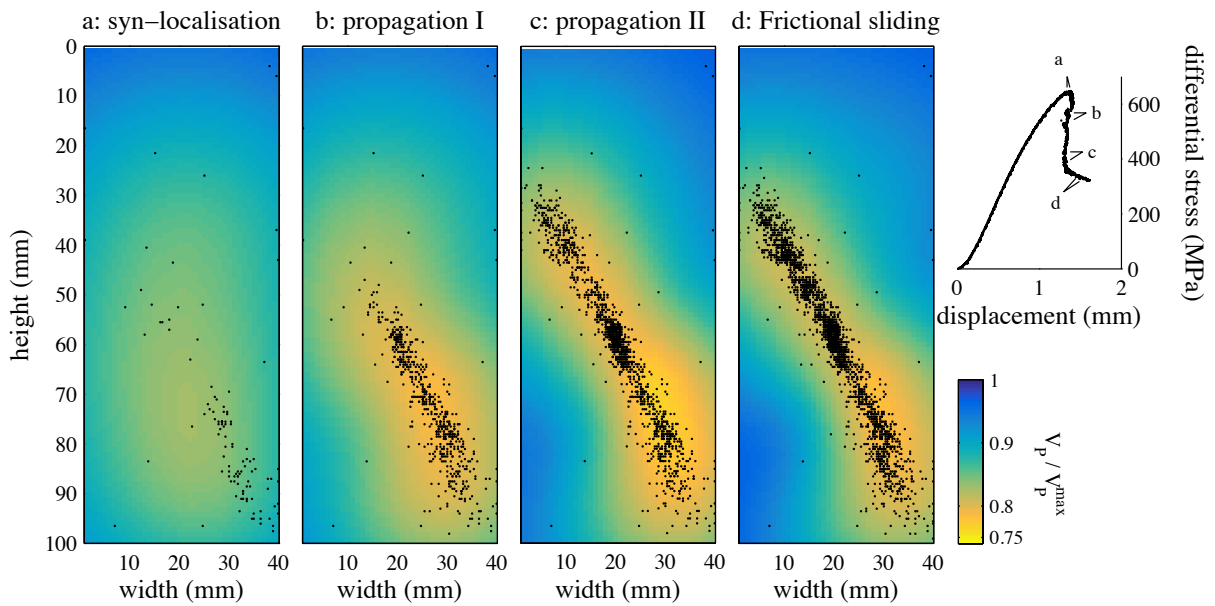


Figure 5: Tomographic sections of the horizontal V_p normalised to the initial velocity through the centre of sample LN5 (quasi-static rupture propagation), perpendicular to the fault. The four slices show time intervals **(a)** during localisation of deformation, **(b)**, **(c)** during two stages of rupture propagation, and **(d)** during frictional sliding of the fault. The corresponding parts of the stress-displacement curve are indicated on the right. AE source locations up to the time interval shown are projected onto the slice. The AE source locations are within 2.5 mm distance perpendicular to the slide, and were determined using the 3D seismic velocity model. Figure from *Aben et al.* [2019].

the fault zone (Figure 5b, c). V_p within the low velocity zone is as low as 4.6 km/s (a 25% drop relative to unaffected areas outside of the zone). In the wake of the rupture tip, the P -wave velocity at some distance from the fault recovers by at most 5%. After rupture completion at the onset of frictional sliding, V_p recovers throughout the sample (the minimum V_p rises by about 100 m/s, Figure 5d).

Mixed rupture propagation: The horizontal V_p before failure decreases throughout the sample (Figure 6a), similar to the velocity drop observed before dynamic failure and quasi-static rupture (Figures 4a and 5a). The deformation history of this particular sample becomes somewhat complicated after the peak stress: We observe a faint localisation zone, delineated by AE source locations and visible in a polished section (Figure d), that is oblique to the final failure surface. The V_p structures of the first few time intervals after the peak stress show a low velocity zone around this aborted nascent rupture plane (Figure 6b). The eventual fault forms after a 50 MPa drop relative to the peak stress, and is embedded in a zone with velocities as low as 4.6 km/s, from an initial velocity of 5.9 km/s – a 22% drop (Figure 6c). The rupture was allowed to propagate dynamically at about 520 MPa. The V_p structure after failure shows two elongated low velocity zones, one around the main fault zone and one around the ‘failed’ fault zone with velocity reductions of 20% and 18%, respectively, a recovery by 200 m/s up to 4.8 km/s (Figure 6d).

A localised low P -wave velocity zone is observed for the three rupture types (quasi-static, dynamic, and mixed). The minimum horizontal velocities within these zones are of similar order of magnitude: 4.6–4.8 km/s, equal to a 22–25% drop relative to the initial P -wave velocity. These velocity drops are in accordance with the largest drops observed in the path-averaged horizontal V_p (Figure 3). The largest velocity decrease for quasi-static and mixed rupture is observed during the propagation of the rupture itself (Figures 4c and 6c). For dynamic rupture, the lowest velocities were observed directly after failure and thus provide only an upper bound for the lowest horizontal P -wave velocities during dynamic rupture.

3.3.1. P -wave tomography: Anisotropy

During axial loading up to failure, V_p anisotropy in dynamic and mixed rupture samples is fairly homogeneous and varies between 10–11% (i.e., the vertical P -wave velocity is 10–11% higher than the horizontal P -wave speed). Some variation in anisotropy near the sample extremities may be caused by lateral confinement from the coupling with the loading column. The anisotropy for the quasi-statically ruptured sample is somewhat higher at 13–15%, although a 1–2% variance within the sample is similar to the dynamic and mixed rupture samples.

The V_p anisotropy adjacent to the ruptured zone increases up to 20% during quasi-static rupture. The anisotropy outside the ruptured zone remains at 15%, similar to the pre-rupture anisotropy. Anisotropy measured in the mixed rupture sample, during the quasi-static rupture interval, increases to 19% around the ruptured zone, and anisotropy outside the ruptured zone remains more or less constant at 12%. Thus, during rupture the vertical P -wave velocity decreases less relative to the horizontal P -wave velocity.

The lowest anisotropy after rupture completion (at residual shear stress) is observed in the dynamically ruptured sample, with a maximum anisotropy of 14% near the ruptured zone and a 11% anisotropy outside this zone (Figure 7a). The maximum anisotropy near the ruptured zone after completion of quasi-static rupture is 19% (Figure 7b), which is a small recovery relative to the maximum anisotropy during rupture. The anisotropy in the volume unaffected by rupture remains similar to the pre-rupture anisotropy. The anisotropy after dynamic failure in the mixed rupture sample is 20% (Figure 7b), and the minimum anisotropy outside the ruptured zone is 12%. We thus see in all three rupture experiments an increase in anisotropy around the ruptured zone during and after failure, with the smallest increase after dynamic rupture. We can infer from the horizontal P -wave velocity decrease and anisotropy increase in the ruptured zone that the vertical P -wave velocity during rupture does not change much. Outside the ruptured zones, the anisotropy during and after rupture remains constant relative to the initial anisotropy just prior to reaching the peak differential stress.

3.4. Microstructural observations

Study of thin sections by optical microscopy reveals a zone of microfractures of around 1 mm in length and oriented parallel to

the loading direction enveloping the shear failure zone (Figure 8a). For quasi-static shear failure, the extent of this damaged zone is appraised at roughly 8 to 10 mm on the tensile side of the fault, and 2 to 3 mm on the opposite compressive side. Several grains outside this off-fault damage zone have been subjected to extensive fracturing as well (Figure 8a). We will attempt to quantify our qualitatively assessed order of magnitude damage zone width from fracture density data obtained from SEM images hereafter. First, we describe the microstructures observed at smaller scale in the SEM images, followed by measurements of off-fault microfracture orientation and density.

The SEM images show that the main failure plane resulting from dynamic rupture is surrounded by patches of gouge and cataclasite (Figure 8b, c), which were not preserved everywhere in the sample during the post-mortem treatment. Whereas the individual particles in patches of gouge cannot be clearly distinguished on the images, the fragments in the cataclasite zones are clearly visible and angular, and show rotation relative to their neighbouring fragments. At 100–500 μ m distance from the main failure zone, the rock contains abundant mode I microfractures oriented parallel to the main loading direction with little to no shear or rotation of fragments (Figure 8b). Some of these mode I fractures tend to deflect towards the main slip zone (Figure 8b). Qualitatively, the amount of microfractures decreases with increasing distance from the fault (Figure 8b, c; Figure 9a, b), and variation in microfracture density on the scale of individual SEM images is linked to mineral type (for instance, the biotite grain at the top of 9b is more heavily fractured relative to the feldspar below it). The microstructural damage observed near a quasi-statically formed failure zone is qualitatively similar to that observed after dynamic rupture: patches of gouge and cataclasite zones (Figure 8d, e) are visible along the main failure zone and some cataclasite zones form secondary brittle shear zones (Figure 8d). Other parts along the main failure plane are less complex and show only a thin zone of gouge and cataclasite (Figure 8e). Primarily mode I microfracture damage is observed further away from the fault (Figure 8d, e). The scope of this study is to quantify off-fault damage related to rupture, and we therefore removed from the traced images the zones of gouge and cataclasite that are clearly related to slip before further analysis of off-fault microfracture damage. Examples of SEM-BSE images with fracture traces are shown in Figure 9a, b and Figure 10a, b.

Off-fault microfracture orientations: The dominant fracture orientation for off-fault microfractures was obtained from the cumulative length of the major ellipse axis of all the fracture segments in all SEM images that fall within 5° intervals measured relative to the loading axis. All intervals are normalised by the interval with the largest cumulative length. The overall dominant off-fault microfracture orientation is parallel to the loading axis for both dynamic rupture (Figures 9c) and quasi-static rupture (Figure 10c). The angle of the off-fault microfractures with respect to the fault plane is somewhat larger for the dynamic rupture case relative to the quasi-static one.

Off-fault fracture density: The surface area of, and fracture traces in, gouge and cataclasite zones and empty fault space has not been used in the calculation of the off-fault microfracture density. The off-fault microfracture density is presented as a function of fault perpendicular distance. We set the origin of each SEM transect (i.e., 0 mm fault perpendicular distance) at the centre of the main failure plane, whose width varies along the fault but remains less than 1 mm (Figure 8) – thus none of the SEM images is entirely located in the main failure plane.

After dynamic rupture, off-fault fracture density, ρ^{frac} , is around 80 mm/mm² directly adjacent to the failure zone, and drops to 30–40 mm/mm² at about 1 mm distance from the failure zone (Figure 9d). Further from the fault, we observe an overall cm scale trend of decreasing ρ^{frac} with increasing distance from the fault, superimposed to a mm scale variation. This variation is between 5–30 mm/mm² and also decreases with distance from the fault (Figure 9d).

ρ^{frac} is around 50 mm/mm² directly adjacent to the quasi-statically formed failure zone (Figure 10d). Within 1 mm distance, ρ^{frac} drops to about 10–20 mm/mm². After this initially steep

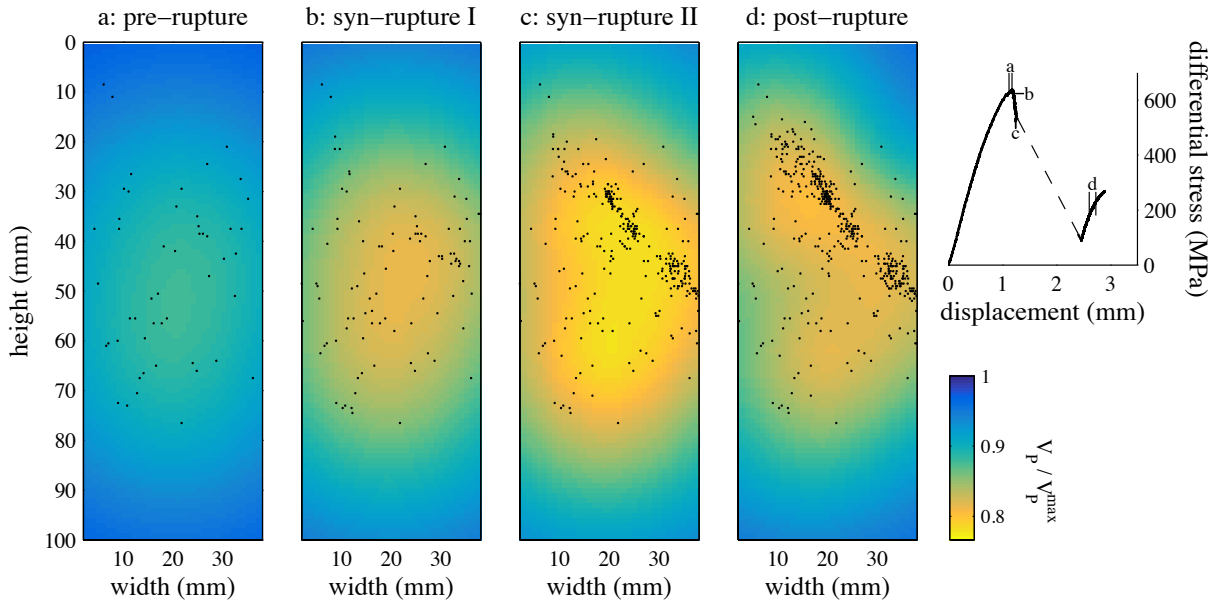


Figure 6: Tomographic sections of the horizontal V_p normalised to the initial velocity through the centre of sample LN8 (mixed rupture propagation), perpendicular to the fault. The four slices show time intervals **(a)** during localisation of deformation, **(b)** and **(c)** during two stages of rupture propagation, and **(d)** during reloading of the fault. The corresponding parts of the stress-displacement curve are indicated on the right. AE source locations up to the time interval shown are projected onto the slice. The AE source locations are within 2.5 mm distance perpendicular to the slide, and were determined using the 3D seismic velocity model.

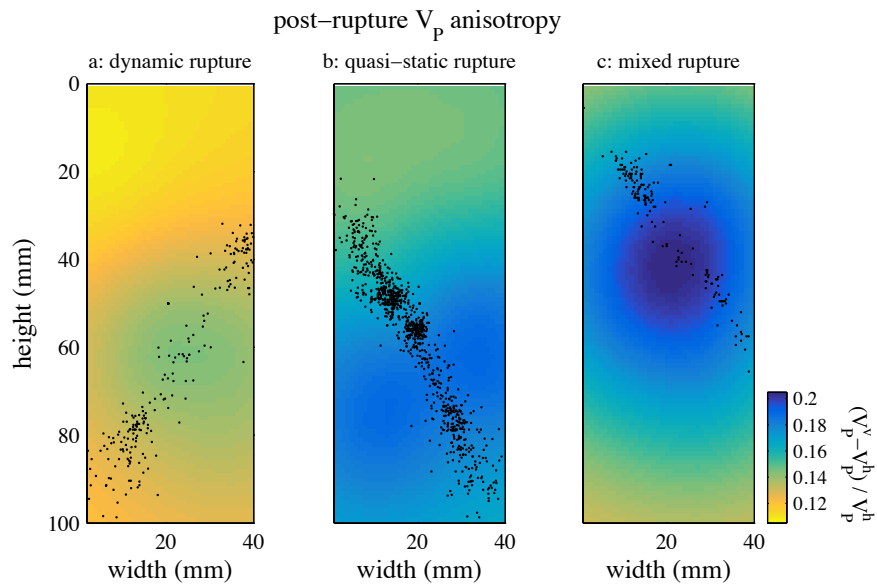


Figure 7: P -wave anisotropy tomography sections through the centre of the samples and oriented perpendicular to the failure planes, based on ultrasonic data measured during reloading of a dynamically failed sample **(a)**, after quasi-static rupture completion **(b)**, and during reloading of a mixed failure experiment **(c)**. The tomography sections in **(a)**, **(b)**, and **(c)** are from the same time intervals as the horizontal P -wave results in Figure 4d, Figure 5d, and Figure 6d, respectively. AE source locations within 5 mm perpendicular to the section are projected on the section, and only post-rupture AE events are shown.

drop, ρ^{frac} decreases to below 10 mm/mm² over 1.5 cm fault perpendicular distance. A mm scale variation is restricted to about 10 mm/mm² magnitude, and decreases with distance. ρ^{frac} measured across the quasi-statically ruptured failure zone is lower and has a lower variance relative to ρ^{frac} measured across the dynamically ruptured failure zone.

3.4.1. Damage zone width

We shall now attempt to summarise the off-fault microfracture density data in an informative and simple measure. For this, we elect the measure of a damage zone width, allowing a direct comparison with fault damage zones studies in the field and with models that predict the extent of off-fault fracture damage from fault rupture and fault slip. The order of magnitude estimate for damage zone width from optical microscopy will be an independent indicator in determining damage zone widths from microfracture density data.

The fracture damage that was measured in the SEM images was created during three stages, according to AE activity: i) pre-failure microfracturing throughout the volume of the sample during yield, leading to ii) fracture coalescence in the nucleation patch and process zone of the propagating rupture, which is followed by iii) slip-induced damage [Tapponnier and Brace, 1976]. Here, we are only interested in microfractures formed during stage ii), as it provides a measure for Γ_{off} . Although we suppress clearly slip-induced stage iii) damage by removing gouge and cataclast layers from the traced images, we cannot rule out that some of the off-fault microfracture damage has a slip-related origin.

The damage zone width is defined as the fault perpendicular distance where the fault-related fracture density trend (stage ii) and iii) damage) intersects the background fracture density. We define the background fracture density ρ_0^{frac} as the sum of yield-related damage and initial damage already present in the samples prior to the experiment. It is reasonable to assume that stage i) introduced equal amounts of damage in the dynamic and quasi-static failed samples, and so ρ_0^{frac} estimated for the quasi-statically ruptured sample represents ρ_0^{frac} of the dynamically ruptured sample as well. We note that in the first place, our definition of background fracture damage gives a reference value for ρ_0^{frac} for our experiments, and is not, but may approach, the background fracture density as encountered in field studies.

We obtain a measure for the background fracture density ρ_0^{frac} for the quasi-static failed sample from the average fracture density of 19 SEM images. The conditions to assume that these SEM images were outside the damage zone were: 1) They are located at 12 mm distance or more from the failure zone, 2) they lack open microfractures more than half the image in length, and 3) they do not contain heavily fractured zones (some example images are shown in SI Text S3). Of these 19 SEM images, the images with the lowest fracture densities (around 2-4 mm/mm²) are qualitatively similar to the initial undeformed state of Lanhélin granite [Siratovich et al., 2105] – but ρ_0^{frac} based on these images excludes yield-related fracture damage and would give a lower bound only. To obtain a more realistic ρ_0^{frac} , the 19 SEM images also include some with a higher background fracture damage (up to 10 mm/mm²), but without clear stage ii) or iii) related fracture damage.

We find that $\rho_0^{\text{frac}} = 6.2$ mm/mm² with a standard error of 2.8 mm/mm² (Figure 9d, 10d). A visual check makes it clear that the trend of decreasing fracture density after quasi-static failure intersects with the background fracture density at around 5 to 10 mm fault parallel distance for most transect (Figure 10d), which matches the damage zone width first estimated from optical microscopy images (Figure 8a). Fracture density data after dynamic failure seems to intersect with the established background fracture density at larger fault perpendicular distances between 10 to 20 mm, suggesting a wider damage zone (Figure 9d).

The decrease in fracture density with distance may be described by a power law function or exponential function, as is often done for damage zone studies in the field and laboratory [e.g., Faulkner et al., 2011; Mitchell and Faulkner, 2009; Savage and Brodsky,

2011; Moore and Lockner, 1995; Ostermeijer et al., 2020]. The intersection of such a fitted function intersects with the background fracture density threshold provides a damage zone width. This approach applied to a high resolution off-fault damage dataset with a large natural variance results in very large uncertainties on the damage zone width [Ostermeijer et al., 2020], thus oversimplifying or misrepresenting the actual damage distribution. We nonetheless pursued this approach for each transect and the results are presented in Text S3. For most transects, both on quasi-statically or dynamically failed samples, a power law decay fits best with the data. The damage zone width results are not always sensible and in harmony with our primary observations (Figure 8a, Figure 9d, and 10d) – for instance, damage zone widths after quasi-static rupture that are much larger than 20 mm (i.e., outside the sample). We therefore use this method merely as an additional guidance, and resolve to manually picking damage zone widths for all transects in both samples. Note that our approach for obtaining damage zone width may differ from that used in field studies: For instance, we may have used a different definition for background fracture density, and we combined additional constraints with the results of fitting a damage decay function.

The damage zone widths determined for the quasi-statically ruptured sample are between 7 and 13 mm on the tensile side of the fault, and between 4 and 13 mm on the compressional side of the fault (Figure 11c). These values are in accordance with our simple estimates from optical microscopy. The damage zone width may exceed the measured transect length or the sample width for a number of transects in the dynamically ruptured sample, in which case we ascribe a lower bound value of 20 mm. On both sides of the fault, the damage zone width is between 11 to 20 mm (Figure 11c). The damage zone after dynamic failure is thus wider by about a factor of two relative to the damage zone created during quasi-static failure. We did not observe a clear trend between damage zone width and distance from rupture nucleation for dynamic and quasi-static rupture (Figure 11c). The power law exponents of the highest quality power law fits vary between -0.37 and -0.49 for transects in both samples, these values are similar to those obtained for fault damage decay profiles in crystalline rock in the field [Savage and Brodsky, 2011; Ostermeijer et al., 2020].

4. Discussion

4.1. Ultrasonics and Tomography

P-wave velocity variations during quasi-static and dynamic failure experiments can be ascribed to two effects: 1) Variations in differential stress, where an increase results in closing of pre-existing horizontal microfractures (i.e., perpendicular to the loading axis) and opening of pre-existing vertical microfractures. Closing of horizontal microfractures mostly affects vertical V_p , and opening of vertical microfractures has a greater effect on the horizontal V_p [Paterson and Wong, 2005, Chapter 5]. 2) Microfracture formation and growth during quasi-static or dynamic failure reduce the *P*-wave velocity locally. These microfractures are subjected to opening or closing as well.

These two effects are recognised during and after our shear failure experiments, where predominantly vertical microfractures are formed within the damage zone (Figures 9c and 10c). These fractures decrease path-averaged *P*-wave velocities along horizontal ray paths more than those along angled ray paths, causing the observed anisotropy (Figure 3). The closing of the vertical microfractures, caused by the syn-failure differential stress drop, results in partial *P*-wave recovery [Passelegue et al., 2018]. The lowest V_p along ray paths intersecting the failure zone are thus observed when the contribution of fracture opening in the rupture process zone dominates over the contribution of fracture closure due to decreasing differential stress (Figure 3a, asterisks). We see a similar evolution in V_p in the time-resolved 3D *P*-wave structure during the stress drop associated to quasi-static failure (Figure 5b, c): A recovery of V_p throughout the sample, except near the rupture front where V_p decreases.

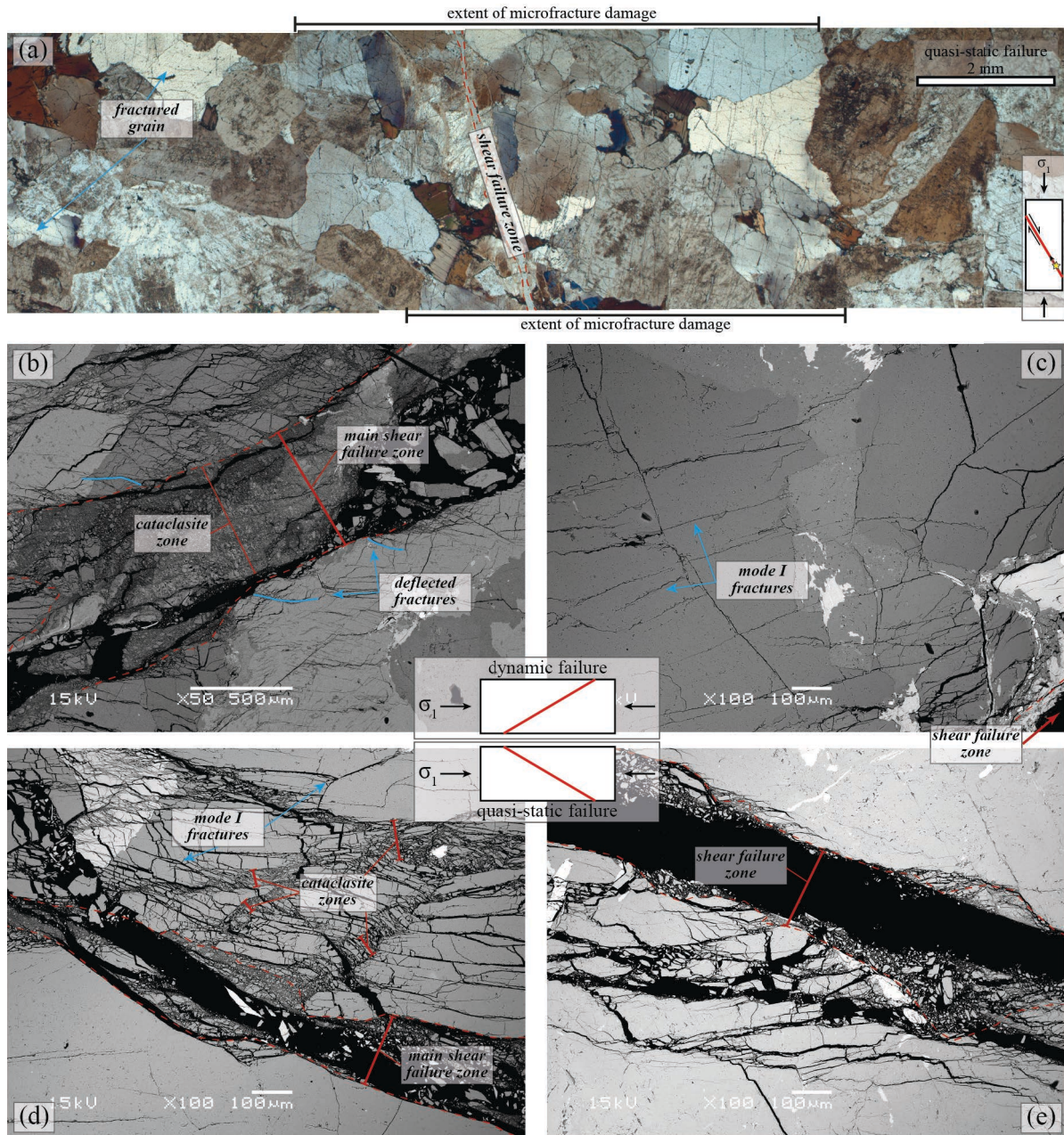


Figure 8: **(a)**: Optical microscope images (transmitted light, cross polarizers at 45°) perpendicular to the failure zone after quasi-static failure show a damage zone of mm-scale microfracture damage, and heavily fractured individual grains outside this zone. Star and arrow in the inset indicates rupture nucleation and propagation direction. SEM-BSE images of **(b),(c)**: Dynamically (sample LN7) and **(d),(e)**: Quasi-statically (sample LN5) failed samples. The principal loading axis is oriented horizontally in all four images. **(b)**: The main failure plane is delineated with gouge and in some places bound by zones of micro-cataclasite. Surrounding grains show mode I microfractures, of which some deflect into the slip zone. **(c)**: Off-fault mode I microfracture damage within an individual grain decreases somewhat with distance from the failure plane (bottom right corner). **(d),(e)**: Secondary shear deformation zones (cataclasite zone) branch of the main failure plane (d), whereas other sections of the main failure plane are relatively straight without secondary structures (e). Off-fault mode I microfracture damage is visible in all cases, but is higher where there is more complexity on the main failure plane.

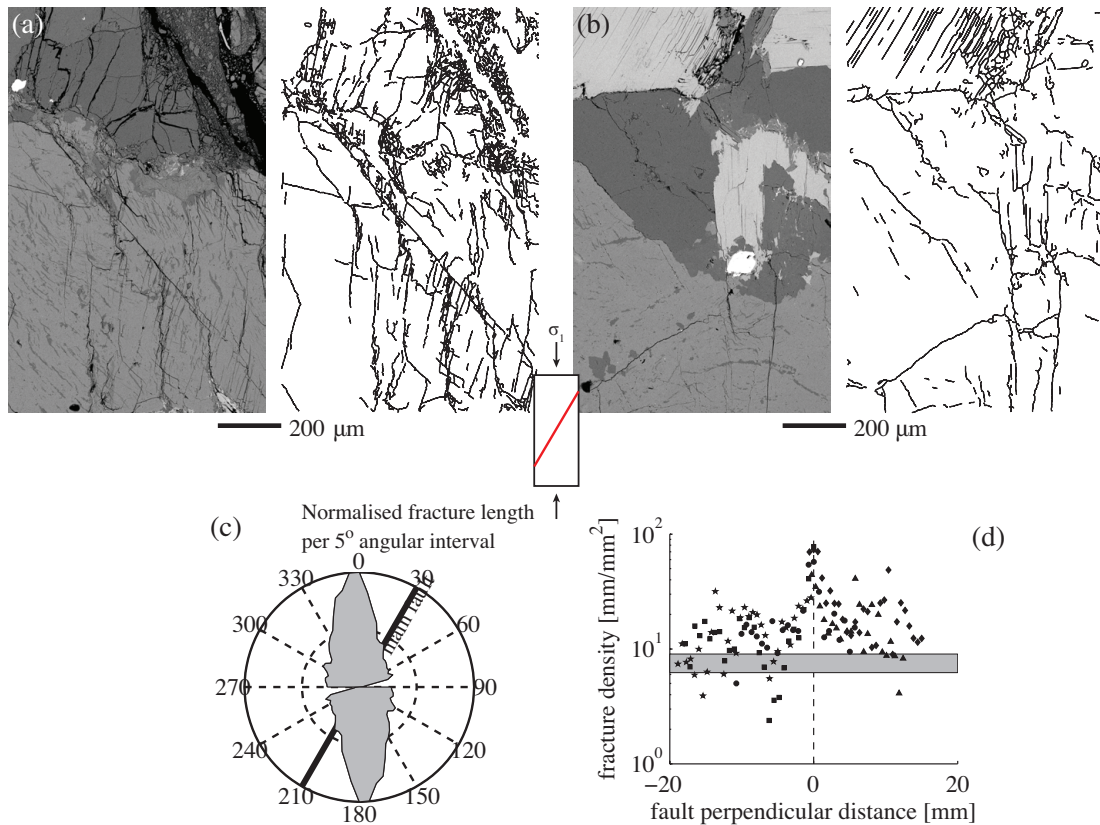


Figure 9: Microstructural analysis of a dynamically ruptured sample (sample LN7). **(a)**: Left panel: SEM-BSE image (100× magnification) near the main fault. The main fault is visible at the top right, surrounded by abundant fault gouge. Right panel: Fracture traces obtained from the images on the right. **(b)**: SEM-BSE image (left panel) and fracture traces (right panel) at some distance from the main fault. The locations of (a) and (b) within the sample are indicated in Figure 1a. **(c)**: Cumulative fracture length per 5° angular interval for all SEM images indicates predominantly sub-axial microfracturing. Cumulative length has been normalised by the largest cumulative length. **(d)**: Fracture density per image as a function of fault-perpendicular distance. The different symbols indicate the different transects defined in Figure 1a, and the gray area bounds the background fracture density range between the mean and the mean plus one standard error that was established on the quasi-statically ruptured sample LN5.

After dynamic failure, at the onset of reloading, the path-averaged V_p along most ray paths rises by a few percent. This is followed by a slight decrease for the remainder of the reloading interval (Figure 3b). In the 3D velocity structure, the post-rupture V_p initially increases, in particular within the low velocity zone (Figure 4b, c). V_p further decreases with progressive reloading (Figure 4c, d). This suggests that horizontal microfractures are closed immediately after the dynamic stress drop, resulting in a P -wave velocity increase. After horizontal microfracture closure, the opening of vertical microfractures dominates and horizontal P -wave velocity decreases. Such a progression is typically observed at the onset of loading of crystalline rock [Paterson and Wong, 2005, Chapter 5].

The path-averaged P -wave velocities after quasi-static failure are very similar in magnitude to the P -wave velocity measured along the same ray paths after dynamic failure and reloading (Fig-

ure 3). The V_p drop during reloading of the dynamically failed sample is a near perfect extension of the V_p increase during the transition to frictional sliding – the V_p -stress curves of each pair of matching ray paths can be connected fairly well, except for ray path E. This horizontal ray path crosses the fault zone and shows a much larger velocity drop near 300 MPa differential stress in the dynamic case than it does near the same differential stress for the quasi-static case. This may reflect the wider damage zone that was created during dynamic rupture, as this wave path is more sensitive to vertically oriented micro fracture damage than the other wave paths that intersect the fault.

4.2. Relationship between microfracture damage and physical properties

P -wave velocity and anisotropy changes are a direct result from changes in the effective elastic moduli of the material. Under the

assumption that effective elastic moduli changes are primarily induced by the formation of microfractures, the P -wave tomography data contain information about the microfracture density. We use an effective medium approach to relate the seismic velocity from the tomographic data to effective elastic moduli, and obtain a fracture density tensor from these effective elastic moduli. We then compare the obtained microfracture density tensor with the microfracture densities measured on thin sections.

4.2.1. Fracture density computation following an effective medium approach

We adopt the effective medium approach by *Sayers and Kachanov* [1995] for a solid containing non-interacting penny-shaped cracks. The overall strain in a cracked solid is the sum of the elastic strain in the matrix (i.e., the constituent minerals of the rock) and the additional strain due to the presence of cracks:

$$\varepsilon_{ij} = S_{ijkl}^0 \sigma_{kl} + \Delta S_{ijkl} \sigma_{kl}, \quad (1)$$

where is S_{ijkl}^0 the elastic compliance tensor of the matrix, σ_{kl} is the stress tensor, and ΔS_{ijkl} is the change in elastic compliance resulting from cracks given by [*Sayers and Kachanov*, 1995]:

$$\Delta S_{ijkl} = \frac{1}{4} (\delta_{ik} \alpha_{jl} + \delta_{il} \alpha_{jk} + \delta_{jk} \alpha_{il} + \delta_{ji} \alpha_{ik}) + \beta_{ijkl}. \quad (2)$$

Here, δ_{ij} is the Kronecker delta, and

$$\alpha_{ij} = \frac{32(1 - \nu_0^2)}{3E_0(2 - \nu_0)} \frac{1}{V} \sum_r (a_r)^3 n_i^r n_j^r \quad (3)$$

is the second rank crack density tensor for r penny-shaped cracks with radii a_r and unit normal vector n_i^r in a volume of rock V with intact matrix elastic parameters E_0 and ν_0 (Young's modulus and Poisson's ratio, respectively). β_{ijkl} is a fourth rank crack density tensor, the contribution of which can be neglected in case of a dry rock with low Poisson's ratio [*Sayers and Kachanov*, 1995].

The tomographic inversion algorithm for the P -wave velocity allows for a vertical transverse isotropy in each voxel, which would result from a transversely isotropic orientation distribution of cracks. The orientation distribution of off-fault fracture segments observed across the quasi-static and dynamically formed failure zones (Figures 9c and 10c) shows a dominant orientation that is near-vertical to the loading axis. We did not measure fracture orientations parallel to the main failure zone, but we assume these are similar to those measured perpendicular to the failure zone so that the microstructural data is consistent with the tomographic models. For a vertical transverse isotropic distribution of cracks, $\alpha_{11} = \alpha_{22}$ are the horizontal components and α_{33} the vertical component of the crack density tensor. *Sayers and Kachanov* [1995] gives the elastic stiffness tensor C_{ijkl} for vertical transverse isotropy in Voigt notation as:

$$\begin{aligned} C_{11} + C_{12} &= (S_{11}^0 + \alpha_{33})/D, \\ C_{11} - C_{12} &= 1/(S_{11}^0 - S_{12}^0 + \alpha_{11}), \\ C_{33} &= (S_{11}^0 + S_{12}^0 + \alpha_{11})/D, \\ C_{44} &= 1/(2S_{11}^0 - 2S_{12}^0 + \alpha_{11} + \alpha_{33}), \\ C_{13} &= -(S_{12}^0)/D, \\ C_{66} &= 1/(2S_{11}^0 - 2S_{12}^0 + 2\alpha_{11}), \\ D &= (S_{11}^0 + \alpha_{33})(S_{11}^0 + S_{12}^0 + \alpha_{11}) - 2(S_{12}^0)^2. \end{aligned} \quad (4)$$

Within the frame of reference of the sample, the vertical and horizontal crack densities are:

$$\begin{aligned} \rho_v &= \frac{2\alpha_{11}}{h}, \\ \rho_h &= \frac{\alpha_{33}}{h}, \\ h &= \frac{32(1 - \nu_0^2)}{3E_0(2 - \nu_0)}. \end{aligned} \quad (5)$$

To find values for ρ_v and ρ_h , we use a similar inversion protocol to [*Brantut*, 2015] where we calculate the theoretical compliance tensor C_{ij} for a range of possible values of ρ_v and ρ_h . From C_{ij} , we obtain synthetic values for V_P^v and V_P^h :

$$\begin{aligned} V_P(\theta) &= \frac{C_{11} \sin^2(\theta) + C_{33} \cos^2(\theta) + C_{44} + \sqrt{M}}{2\rho} \\ M &= \left[(C_{11} - C_{44}) \sin^2(\theta) - (C_{33} - C_{44}) \cos^2(\theta) \right]^2 + [(C_{13} + C_{44}) \sin^2(\theta) + (C_{33} - C_{44}) \cos^2(\theta)]^2 \end{aligned} \quad (6)$$

where ρ is the density of the intact rock matrix and θ the angle with respect to the loading axis, which is 0° for V_P^v and 90° for V_P^h . We then use a least absolute criterion to obtain the best fit between synthetic V_P and measured V_P , assuming a Laplacian probability density function, so that we obtain the most likely values for ρ_v and ρ_h [*Tarantola*, 2005; *Brantut et al.*, 2011]. For this, we assume an uncertainty on the measured V_P of 200 ms^{-1} .

The P -wave tomography models obtained after quasi-static, dynamic, and mixed failure provide the observed values for V_P^h and V_P^v . We took these velocities along a fault-perpendicular transect through the centre of each sample. S_{ij}^0 and h were calculated from $\nu_0 = 0.20$ and E_0 . We used a value for E_0 derived from the path-averaged V_P measured at peak stress, and take $\rho = 2660 \text{ kg m}^{-3}$.

Computed vertical crack densities after dynamic failure increase from $\rho_v = 0.09 - 0.10$ at the edge of the sample to $\rho_v = 0.18$ near the failure zone (Figure 12a). The horizontal crack density ρ_h increases from 0 near the edge of the sample to about 0.04 near the failure zone. These increasing crack densities suggest a damage zone width of about 20 mm on both sides of the fault, but we exercise caution with this measure as it is near the resolution of the tomography imposed by the correlation length. After quasi-static failure, computed crack densities near the edge of the sample are somewhat lower compared to those computed for dynamic failure ($\rho_v = 0.04 - 0.07$ and ρ_h is negative, Figure 12b), but show a stronger increase near the failure zone where they show the same peak values ($\rho_v = 0.18$ and $\rho_h = 0.03$). The negative fracture densities near the edge of the sample result from a slight underestimate of the value for E_0 , which was obtained from path-averaged V_P measurements. Such an error is expected in the absolute values of the path-averaged V_P , but has minor consequences for the change in crack densities. The crack densities after mixed failure are similar to those computed after quasi-static and dynamic failure, but show a strongly asymmetric distribution across the failure zone (Figure 12c). The higher crack densities on one side of the fault (positive fault-perpendicular distance in Figure 12c) coincide with the nascent secondary failure zone (Figure 2d).

4.2.2. Off-fault microfracture density from microstructures compared to crack density

The horizontal and vertical crack densities, ρ_h and ρ_v , computed from ultrasonic wave velocities have units of m^3/m^3 and are directly derived from the crack tensor components α_{11}/h and α_{33}/h (equation [5]). The off-fault microfracture density ρ^{frac} obtained from microstructures is measured in m/m^2 and is a scalar quantity. In order to compare the two methods, we convert the off-fault microfracture traces to tensor components α_{11}/h and α_{33}/h . We remain in the spirit of the effective medium approach by assuming that the sample contains a transversely isotropic orientation distribution of penny-shaped fractures so that $\alpha_{11} = \alpha_{22}$ and all cracks have a radius a_r . We treat each traced fracture segment as a trace through an individual penny-shaped fracture.

The centre of a traced fracture does not necessarily lie on the SEM image plane, which means that the fracture trace length t is equal to or smaller than the true fracture diameter $2a$. The mean fracture radius \bar{a} was obtained from the mean measured trace length \bar{t} so that $2\bar{a} = (\pi/2)\bar{t}$ [*Oda*, 1983]. To determine \bar{t} , we first determined the probability distribution of trace lengths. We followed the

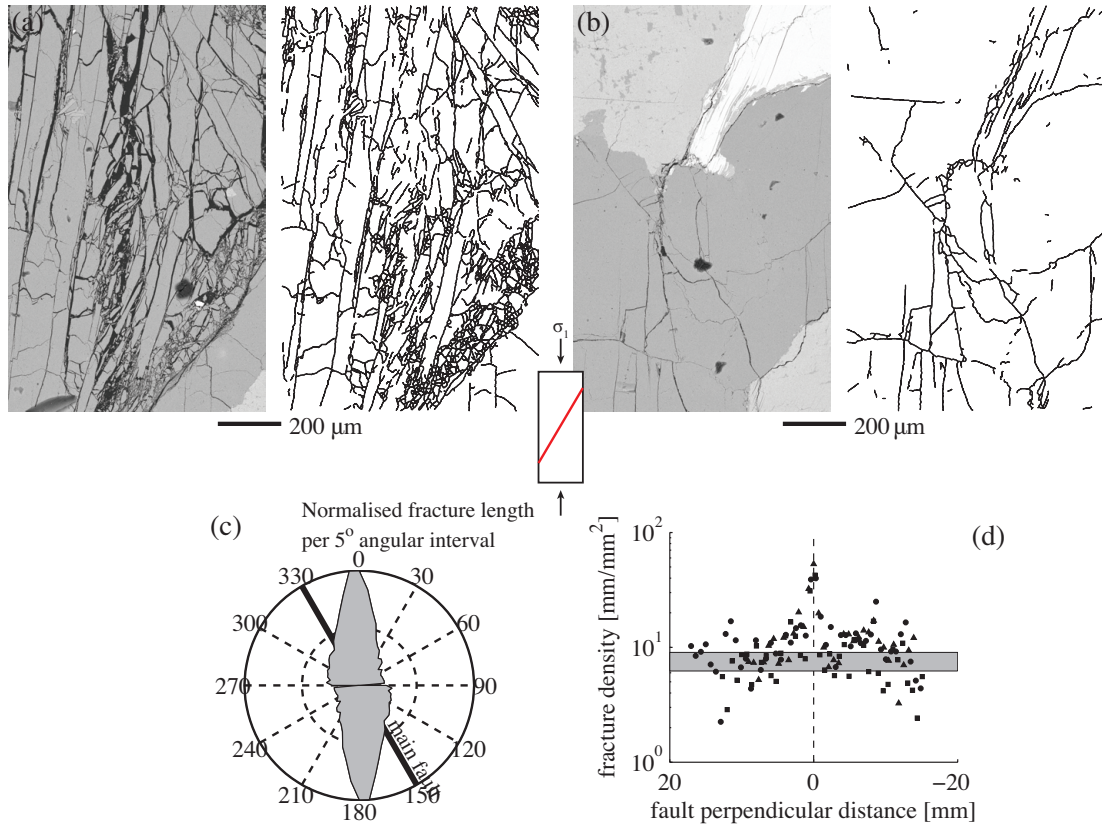


Figure 10: Microstructural analysis of a quasi-statically ruptured sample (sample LN5). **(a)**: Left panel: SEM-BSE image (100× magnification) near the main fault. The main fault is visible at the bottom right, surrounded by abundant fault gouge. Right panel: Fracture traces obtained from the images on the right. **(b)**: SEM-BSE image (left panel) and fracture traces (right panel) at some distance from the main fault. The locations of (a) and (b) within the sample are indicated in Figure 1a. **(c)**: Cumulative fracture length per 5° angular interval for all SEM images indicates predominantly sub-axial microfracturing. Cumulative length has been normalised by the largest cumulative length. **(d)**: Fracture density per image as a function of fault-perpendicular distance. The different symbols indicate the different transects defined in Figure 1a, and the gray area bounds the background fracture density range between the mean and the mean plus one standard error.

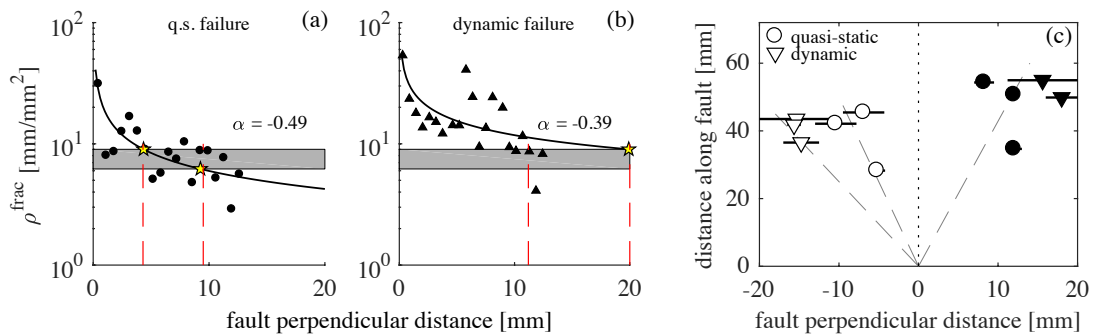


Figure 11: Fracture density along a single fault perpendicular transect after quasi-static **(a)** and dynamic **(b)** failure, fitted with a power law function (exponent α given for both fits). Yellow stars indicate the intersection of the fitted function with the gray area, which bounds the background fracture density range between the mean and the mean plus one standard error. The manually picked range for the damage zone width is shown by the red lines (upper and lower bound). See SI Text S3 for all transects. **(c)**: Distance along the fault versus damage zone width, for transects in samples subjected to quasi-static (circles) and dynamic (triangles) loading. White symbols are transects on the tensile side of the fault, black symbols on the compressional side of the fault (see Figure 1a, b). Distance along the fault is measured in the direction of rupture propagation. Dashed lines indicate the approximate trend.

approach of *Rizzo et al.* [2017] to find the best type of probability distribution that describes the trace lengths in a single SEM image:

1) Maximum likelihood estimators were used to fit power law, ex-

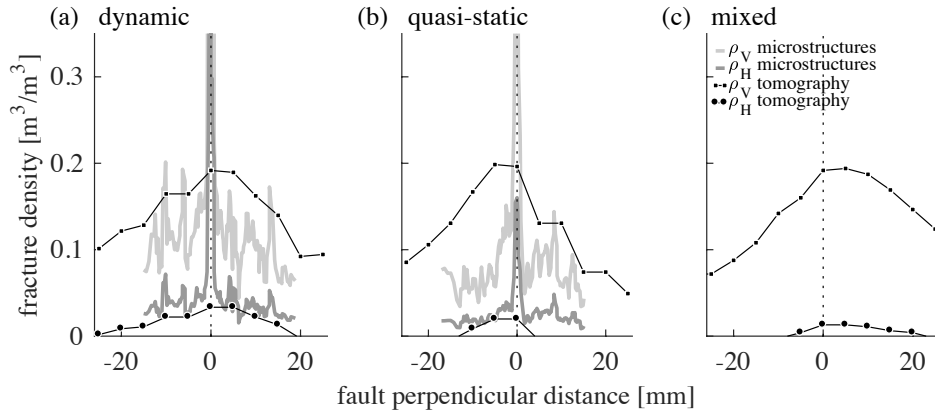


Figure 12: Horizontal and vertical fracture densities (ρ_h and ρ_v) computed from the P -wave tomography models after (a): Dynamic failure, (b): Quasi-static failure, and (c): Mixed failure. The gray curves correspond to horizontal and vertical fracture densities obtained from the off-fault microfracture traces on post-mortem thin section after dynamic and quasi-static failure.

ponential, and log-normal distributions, and 2) the goodness-of-fit for all three types of distributions was tested using the Kolmogorov-Smirnov test, giving a probability for each distribution. We obtained a $> 90\%$ probability for a log-normal distribution of trace lengths for a majority of the images, and a power law distribution for the remaining images. For log-normal distributed trace lengths, \bar{l} was calculated from the first moment of the distribution. \bar{l} cannot be determined from the first moment of a power law distribution, and we therefore took the mean of the measured traces.

The unit vectors in equation [3] for each individual fracture segment are given by $\cos(\theta)$ and $\sin(\theta)$ for the tensor components α_{11} and α_{33} , respectively (Figure 1c). Each image intersects only those fractures that have their centre within \bar{a} distance perpendicular to the image plane [Oda, 1983], under the assumption the average out-of-plane fracture orientations (i.e., rotation with respect to the sample axis) are perpendicular to the image. This assumption agrees with the assumption of a transversely isotropic fracture orientation distribution. The volume V associated to the fracture traces on each SEM image is then given as $V = S\bar{a}$, where S is the surface area of the image. With the parameters \bar{a} , V , and θ , and equations [3] and [5], we obtain ρ_h and ρ_v for each image.

Values for ρ_h and ρ_v obtained from the microstructures agree well with those computed from the tomography models, for both the dynamically and quasi-statically failed samples (Figure 12a, b), except near the failure zone. Here, between 0 and 2 mm distance from the fault, the microfracture densities from microstructures are up to an order of magnitude higher than those obtained from V_p . This can be ascribed to the difference in spatial resolution of the two methods. Nonetheless, the primary (i.e., cm scale) features of the damage around the failure zone are captured by both direct observation of microfractures and by P -wave tomography.

Our results show that P -wave tomography combined with an effective medium theory can quantify localised zones of fracture damage. The use of a normalised fracture density, such as the one presented here, has the advantage of direct applicability with other effective medium models, for instance to predict hydraulic properties [Gavrilenko and Guéguen, 1989; Guéguen and Schubnel, 2003]. We therefore propose that high resolution geophysical measurements of wave speeds from dense arrays [Ben-Zion et al., 2015] combined with microstructural characteristics measured in the field [Rempe et al., 2013, 2018] or from borehole data [Jeppson et al., 2010] can reveal the physical properties around fault zones. Such data can be used to calibrate the findings of rupture simulations that allow for off-fault energy dissipation [Bhat et al., 2012; Thomas and Bhat, 2018; Okubo et al., 2019], and can be compared to laboratory failure experiments such as those presented here.

4.3. Rupture energetics

We first provide estimates for Γ and W_b for all shear failure experiments from the mechanical stress and strain data. We then show that the damage zone width established after quasi-static and dy-

namic failure is the results of stresses induced by rupture and not by slip, so that we can calculate Γ_{off} thereafter. We provide estimates for Γ_{off} for quasi-static and dynamic ruptures based on microstructural observations, and discuss the implications. A similar estimate may be obtained for Γ dissipated on the fault by quantifying the cumulative fracture surface in gouge and cataclasites that make up the main shear failure zone, but this is beyond the scope of this study. We leave such an endeavour to future studies, as difficulties need to be tackled regarding gouge preservation and resolution limits on identifying the smallest gouge grain sizes.

4.3.1. Fracture energy and breakdown work from mechanical data

We calculated breakdown work W_b by converting the measured stress and axial strain data to shear stress and slip along the failure zone, following the steps described by Wong [1982, 1986]. The area under the shear stress versus slip curve in excess of the residual shear stress gives a measure for W_b . We measured a residual shear stress of 140 MPa at the end of the quasi-static rupture experiment after 0.83 mm slip (Figure 2b), whereas the residual shear stress measured after reloading the samples after dynamic and mixed failure was 120 MPa (Figure 2b). This suggests that the quasi-statically created failure zone had not yet reached its residual frictional strength yet, supported by the convergence of the quasi-static failure stress-strain curve towards this value. Residual shear stresses of 140 MPa and 120 MPa after a slip distance of 0.83 mm give us quasi-static values for W_b of 37 kJm^{-2} and 53 kJm^{-2} , respectively. These are lower bounds for quasi-static W_b , as more slip would have been accrued towards continued weakening down to 120 MPa.

Γ for quasi-static failure may be calculated in the same manner, up to the shear stress and slip distance at which the failure zone through the sample was completed. We thereby assume that all breakdown work done to drop the strength of the failure zone from the peak stress down to shear stress of rupture completion was dissipated to create the failure zone, including formation of off-fault microfractures and gouge associated to the rupture. The rupture was completed and the failure zone fully formed at a shear stress of 155 MPa and a slip distance of 0.44 mm, as established by Aben et al. [2019] using AE source locations. From the shear stress and slip data we then obtain $\Gamma = 27 \text{ kJm}^{-2}$ for quasi-static failure [Aben et al., 2019].

W_b cannot be established directly from the mechanical data measured during dynamic failure, as the elastic unloading of the loading column is measured rather than the drop in shear stress of the fault. Order of magnitude estimates for dynamic W_b and Γ may be obtained by approximating the loading system as a simple spring-slider model, similar to Beeler [2001], and solve the force balance

by assuming some slip-weakening law. We tried this approach, but found results that are too erroneous to be useful. This is most likely because the model assumes a constant piston mass during dynamic failure, which is violated at short failure time scales by inertia of the piston.

Lockner et al. [2001] pointed out that it is not strictly correct to use the average shear stress and average slip measured on the sample during failure for calculating W_b (and Γ), since the size of the rupture tip process zone is smaller than the sample size. The assumption in using the average shear stress and average slip for the analysis of W_b [*Rice and Rudnicki*, 1980; *Wong*, 1986] is that the fault is created in the entire sample at the peak stress (i.e., the sample is a point on the trajectory of a propagating rupture) – which we show is not the case. It is however encouraging that by using this approach, similar order of magnitude values for W_b have been found on granitic rock samples with different diameters: 16 mm [*Wong*, 1982], 40 mm [*Aben et al.*, 2019], and 76 mm [*Lockner et al.*, 2001].

4.3.2. Damage zone width

Microfracture damage observed after quasi-static and dynamic failure were induced by both rupture and slip. We observe a wider damage zone after dynamic rupture relative to that observed after quasi-static rupture. Part of the slip-related damage was corrected for by removing gouge- and cataclasis patches from the traced images prior to establishing off-fault microfracture densities and damage zone widths (Figure 11). Nonetheless, some of the remaining off-fault microfractures may be induced by slip rather than rupture. The sample subjected to dynamic failure (sample LN7) accumulated 3.22 mm slip, whereas the sample subjected to quasi-static failure (sample LN5) accumulated 0.83 mm slip. The microstructural record after dynamic failure may thus contain more slip-related off-fault microfracture damage. Before we provide an estimate for $\Gamma_{\text{off}}^{\text{surf}}$, achieved by combining the damage zone width and microfracture density, we assess whether the difference in damage zone width obtained for quasi-static and dynamic rupture is an effect of rupture velocity or an effect of the difference in accumulated fault slip.

Off-fault damage during rupture can be caused by stresses around the rupture tip. The geometry of the rupture tip stress field changes with rupture velocity so that damage is created in a larger area around the rupture tip at higher rupture velocity [*Poliakov et al.*, 2002; *Rice et al.*, 2005], increasing the damage zone width. Slip along a rough fault (i.e., asperities slipping past each other) causes additional stresses in the host rock around the asperities, and these stress heterogeneities result in off-fault damage. With progressive slip along rough faults, progressively larger asperities are dragged past each other and the additional off-fault stresses act over an increasingly larger area [*Chester and Chester*, 2000]. The damage zone width is thus also expected to increase with increasing slip.

The sample subjected to dynamic rupture has experienced both a larger rupture velocity and a larger amount of slip relative to the quasi-statically ruptured sample. Here, we compute the damage zone width as a function of rupture velocity by adopting the analytical solution by *Poliakov et al.* [2002] for the elasto-dynamic stress field in a rupture tip process zone for a non-singular slip-weakening rupture. We make the assumption of small scale yielding: The fracture energy is dissipated before the remainder of the breakdown work is done. This means that the initial drop in shear strength in the rock from peak strength down to 155 MPa is solely ascribed to dissipation of Γ , and further reduction in shear stress is caused by other slip-weakening processes. This assumption seems justified, based on the quasi-static and mixed failure experiments: The initially steep slope in shear stress versus slip during rupture (Figure 2b) causes stronger stress concentrations relative to the less steep slope of the curve after rupture completion. We therefore expect that the initial steep stress drop determines the damage zone width. We also predict the damage zone width as a function of slip by using the analytical solution by *Chester and Chester* [2000] for the stress field along a rough frictional fault in an elastic material. Using these two models, and realistic input parameters obtained

from the rupture experiments, we then assess which parameter (rupture velocity or slip) is responsible for the observed difference in damage zone width between our quasi-static and dynamic rupture experiments.

Rupture tip process zone model: We consider the 2D case of a mode II rupture that propagates parallel to the x -direction at $z = 0$. The stress in the rupture tip process zone is given by:

$$\sigma_{ij} = \sigma_{ij}^0 + \Delta\sigma_{ij}, \quad (7)$$

where σ_{ij}^0 is the far-field stress state on the sample and $\Delta\sigma_{ij}$ are the additional stress components caused by the rupture that are given by equations [A3] and [A11] in *Poliakov et al.* [2002]. To remove the stress singularity at the rupture tip, the shear stress drops linearly from peak stress τ_p to residual strength τ_r over a slip-weakening zone of size R . For an infinite elastic medium, R decreases in size with increasing rupture velocity v [*Rice*, 1980]:

$$R = \frac{R_0}{g(v)} \quad \text{with} \quad R_0 = \frac{9\pi}{32(1-\nu)} \frac{\mu \delta_c}{(\tau_p - \tau_r)}, \quad (8)$$

where R_0 is the quasi-static limit of R , and μ and ν are the shear modulus and Poisson's ratio of the host rock respectively. The function g depends on ν , and on the P - and S -wave velocities of the material [*Poliakov et al.*, 2002]. The model parameters were obtained from the mechanical and ultrasonic data measured during the quasi-static rupture experiment on sample LN5 (Table 2), where we calculated μ from E and ν . These parameters yield $R_0 = 0.14$ m, and R decreases down to 0.02 m at $v = 0.9 \times c_s$. The normal stress on the fault σ_{zz}^0 , stress ratio $k = \sigma_{xx}^0/\sigma_{zz}^0$, and τ_p were calculated from the mechanical data at the onset of rupture. The residual shear stress $\tau_r = 155$ MPa after $\delta_c = 0.44$ mm slip at rupture completion. Note that in the experiments, the shear stress along the failure zone drops further from 155 MPa to 120 MPa, but this occurs after rupture completion and any off-fault damage accrued during this stress drop is not part of off-fault damage related to the rupture.

Rough fault model: We consider the 2D case of a strike-slip fault parallel to the x -direction at $z = 0$. We consider a uniform displacement U along a fault with coefficient of friction μ that has a sinusoidal perturbation:

$$A \sin(2\pi x/L), \quad (9)$$

where L is the wavelength and $A = \gamma L/2\pi$ is the amplitude of the perturbation. γ is a dimensionless roughness factor. The stress around the rough or wavy fault is given by:

$$\sigma_{ij} = \sigma_{ij}^0 + \Delta\sigma_{ij}, \quad (10)$$

where the stress perturbations caused by the sinusoidal fault $\Delta\sigma_{ij}$ are (equations [10a-10c] in *Chester and Chester* [2000]):

$$\begin{aligned} \Delta\sigma_{xx} &= \exp(-lz) [(-1 + lz) \cos(lx) + f(2 - lz) \sin(lx)] B, \\ \Delta\sigma_{zz} &= \exp(-lz) [(-1 - lz) \cos(lx) + (flz) \sin(lx)] B, \\ \Delta\sigma_{xz} &= \exp(-lz) [f(-1 + lz) \cos(lx) + (lz) \sin(lx)] B, \end{aligned} \quad (11)$$

Table 2: Model parameters for the rupture tip process zone model and rough fault model.

	Rupture model	Roughness model
Young's modulus E	88 GPa	44 GPa
Poisson's ratio ν	0.22	0.22
normal stress σ_{zz}^0	262 MPa	180 MPa
stress ratio k	2.2	1.8
fracture energy Γ	27×10^3	-
peak shear stress τ_p	280 MPa	-
residual shear stress τ_r	155 MPa	-
surface roughness γ	-	10^{-2}
coefficient of friction f	-	0.75

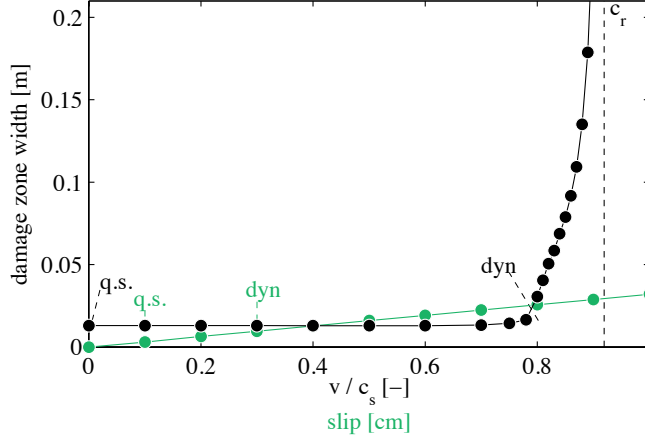


Figure 13: Damage zone width as a function of normalised rupture velocity (black) and fault slip (gray), based on the stresses around a propagating rupture tip [Poliakov *et al.*, 2002] and the stresses caused by a wavy perturbation of a frictional interface [Chester and Chester, 2000]. The rupture velocity of the quasi-static rupture, and the total slip of the quasi-static and dynamic rupture experiments are highlighted. The approximate rupture velocity of the dynamic rupture experiment is indicated where the damage zone width is double that off the quasi-static rupture, in accordance with the microstructural results.

with $l = 2\pi/L$ and

$$B = \frac{\pi E \delta \gamma}{4(1 - \nu^2)}. \quad (12)$$

The normal stress on the fault σ_{zz}^0 , stress ratio $k = \sigma_{xx}^0/\sigma_{zz}^0$, and coefficient of friction f were calculated from the mechanical data at the onset of frictional sliding in the quasi-static experiment (Table 2). The surface roughness was estimated from a post-mortem cross section perpendicular to the fault plane (Figure 2c), where the main fault interface (length of the order of 100 mm) shows a waviness in the order of 1 mm, giving $\gamma = 10^{-2}$. During rupture the elastic constants around the fault interface drop by around 50%, as observed in the P -wave tomography results during and after rupture (Figure 4, 5, and 6). To take into account this elastic weakening by the rupture process zone prior to significant slip, Young's modulus E is half that of the intact rock.

Failure criterion and damage zone width: The 2D off-fault stress tensor around a rupture tip was calculated for a range of rupture velocities from $10^{-6} \times V_S$ (i.e., quasi-static rupture) up to $0.9 \times V_S$ (i.e., rupture velocity near the Rayleigh wave speed). The stress tensor for a wavy fault was calculated for a range of slip distances between 1 mm and 10 mm, for a range of perturbation wavelengths between 0.2 mm and 150 mm. We used a coulomb failure criterion to assess the damage zone width that can be expected from both mechanisms, where the maximum shear stress τ^{\max} and coulomb shear stress τ^{coulomb} are defined as:

$$\begin{aligned} \tau^{\max} &= \sqrt{(\sigma_{xx} - \sigma_{zz})^2/4 + \sigma_{xz}^2} \\ \tau^{\text{coulomb}} &= (\sigma_{xx} - \sigma_{zz}) \sin(\phi)/2 \end{aligned} \quad (13)$$

where $\phi = \tan^{-1}(f^{\text{DZ}})$, and f^{DZ} is the coefficient of friction within the damage zone, for which we take $\mu^{\text{DZ}} = 1$ to represent intact rock [Chester and Chester, 2000]. We expect damage where $\tau^{\max}/\tau^{\text{coulomb}} > 1$ or $\tau^{\max}/\tau^{\text{coulomb}} < 0$. The damage zone width is the largest fault parallel distance at which the failure criterion is satisfied.

The expected fault damage zone width for ruptures propagating below the Rayleigh wave speed increases from about 8 mm at normalised rupture velocities between 0 and 0.75 up to 200 mm or more as the rupture velocity approaches the Rayleigh wave speed (Figure 13). The damage zone width predicted for increasing slip along a rough fault shows a linear increase with slip (Figure 13). The damage zone width observed after quasi-static rupture is around 10 mm (Figure 11c), which is similar to that predicted for a low velocity rupture, whereas the damage zone width predicted to result from slip is less than 3 mm for 0.8 mm of slip accumulated during the experiment. Dynamic rupture resulted in 3 mm total slip, giving a damage zone width of 8.6 mm according to the wavy fault model (Figure 13). This does not match with the observed damage zone width of around 15 to 20 mm (Figure 11c). Although the rupture velocity for this experiment was not measured, a damage zone width of 20 mm can result from a dynamic rupture velocity of about $0.8 \times V_S$.

These results are first order estimates only: The parameters for the rupture model were taken from the quasi-static rupture data, whereas for the dynamic rupture the stress drop may be larger, the coefficient of friction of the fault may be lower, and the breakdown work larger. The rupture model strictly applies to an infinite medium, which may explain why the calculated process zone size R_0 is larger than the actual sample size. For the stresses resulting from slip on a rough fault, the background stresses are assumed to be constant, whereas in our experiments initial slip is accumulated within the rupture process zone where the shear stress and the coefficient of friction are not constant. However, this slip-weakening distance for the quasi-static case is less than 1 mm, after which the applied stresses in the experiment remain more or less constant. The simulated off-fault damage at lower velocity ruptures (< 0.7 normalised velocity) is mostly on the tensile side of the fault plane, similar to the results of Poliakov *et al.* [2002]. Other simulations for rupture-induced off-fault damage also predict a strong asymmetry in off-fault damage distribution, with most damage occurring on the tensile side of the fault [e.g., Rice *et al.*, 2005; Xu *et al.*, 2015; Thomas and Bhat, 2018]. In our experiments however, fracture damage occurs in equal amounts on both sides of the

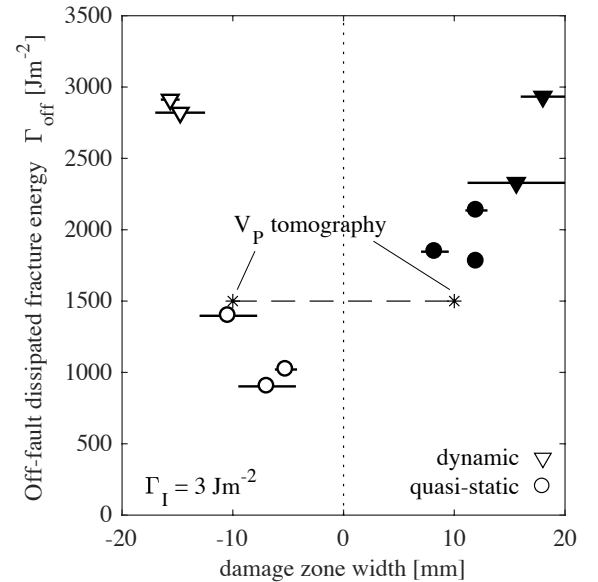


Figure 14: Γ_{off} (calculated using a specific surface energy of 3 Jm^{-2}) versus damage zone width. White symbols are transects on the tensile side of the fault, black symbols on the compressional side of the fault. The asterisks show the damage zone width and Γ_{off} during quasi-static rupture, obtained from P -wave tomography by [Aben *et al.*, 2019].

fault, and some microstructural studies on a off-fault damage surrounding an experimentally formed shear fracture did not observe a clear damage asymmetry either [Moore and Lockner, 1995; Zang et al., 2000]. This may be due to several reasons: 1) Rupture simulations are performed in a large continuum with constant far-field stresses, whereas our experiments were performed on a 100 mm by 40 mm cylinder where boundary effects may alter the off-fault stress fields as described by the models. 2) The orientation of the principal stresses (i.e., stress ratio k) change during failure, which may change the region where the off-fault failure criterion is satisfied [Poliakov et al., 2002; Rice et al., 2005]. In the simulations, the stress ratio k was kept constant whereas it actually changed from $k = 2.2$ to $k = 1.8$ during quasi-static rupture. 3) The trajectory of the propagating rupture is not linear so that the principal stresses with respect to the trajectory of the rupture process zone change locally. 4) The applied stresses change orientation due to already formed fracture damage in and behind the rupture process zone [Faulkner et al., 2006].

Some of the reasons above may be alleviated with a different loading geometry for better control on the principal stress orientations, using for instance a direct shear setup. More advanced rupture simulations with a ‘rough’ rupture trajectory may yield additional insights into the lack of damage asymmetry. Nonetheless, the models suggest that the damage zone width for quasi-static and dynamic rupture in our experiments is controlled primarily by the rupture tip process zone. This is supported by the 3D P -wave velocity structure of the mixed rupture, where a low velocity zone is associated with the incipient secondary fault (Figure 6). This added structural complexity provides a unique opportunity to compare damage in a fault zone without slip, with that in a fault zone that has accumulated 2.4 mm of slip (the main fault in the same sample). The lowest V_P in the zone around the incipient fault is only slightly higher than the lowest V_P around the fully developed fault (Figure 6d), suggesting that rupture rather than slip caused the most off-fault damage.

4.3.3. Estimates for Γ_{off}

A measure for Γ_{off} has been obtained in a previous study from *in situ* V_P tomography measurements on quasi-statically ruptured sample LN5 [Aben et al., 2019]. We now use a second and independent method to obtain Γ_{off} from microstructural data for the same sample, and for a dynamically ruptured sample. The energy dissipated by creating new fracture surface in the volume around the fault gives an estimate for the off-fault dissipated fracture energy Γ_{off} . We assume that the microfractures are mostly tensile – little to no slip and some opening of the microfractures observed in the thin sections testify to this – so that the energy needed to form them is the mode I fracture energy. The cumulative mode I fracture energy gives us Γ_{off} , which was calculated from the fracture density data as follows:

$$\Gamma_{\text{off}} = 2 \sum_{i=1}^n \left(\rho_i^{\text{frac}} - \rho_0^{\text{frac}} \right) x_i \Gamma_1, \quad (14)$$

where x_i is the fault-perpendicular width of image i and Γ_1 is the mode I fracture energy for quartz and feldspar, ranging from 2 to 10 Jm^{-2} [Atkinson, 1987]. A factor 2 is included to account for the two new surfaces that comprise each fracture. We calculated Γ_{off} for all the fracture density transects obtained on both the quasi-statically ruptured and dynamically ruptured samples, for $\Gamma_1 = 2$ and $\Gamma_1 = 10 \text{ Jm}^{-2}$. Γ_{off} averaged for all transects for a quasi-static rupture ranges between 2 and 10 kJm^{-2} , that for a dynamic rupture is between 3 and 15 kJm^{-2} (Figure 14). Γ_{off} for dynamic failure may be higher, considering that the fault damage zone width in the dynamically failed samples may be a lower bound only. Γ_{off} increases nearly linearly with damage zone width, based on the transects through the tensile side of the fault (Figure 14). For the quasi-static case, Γ_{off} on the tensile side of the fault is less than on the compressional side of the fault.

The range of values for Γ_{off} established by the microstructural approach depend mainly on the mode I fracture energy, for which we use values that vary by nearly an order of magnitude (2 and

10 kJm^{-2}), but are usually expected to be at the lower end of these value. We see a good agreement between the results from two independent methods to determine Γ_{off} : Γ_{off} from P -wave tomography falls well within the estimated range for Γ_{off} from microstructures, and is similar to it when the mode I fracture energy is 3 Jm^{-2} (Figure 11c).

We note that not all of the fracture surface area may have been traced from the SEM images, as we elected not to trace fractures below 9 μm since this would have increased noise (for instance, intrinsic flaws in the grains and artefacts from thin section preparation) more than increased actual fracture surface area. A close inspection of the SEM images (Figure 8b-e) reveals that individual fractures shorter than 9 μm nearly all reside in zones of cataclasis and gouge close to the main failure zone, which have been excluded from further analysis. The microfractures further away from the fault zone are generally longer than 30 μm , and so the proportion of short microfractures not included as fracture surface area is small. This is confirmed by the good agreement between the two independent measures for Γ_{off} at a realistic value for the mode I fracture energy.

Γ_{off} measured after dynamic failure is about 1.5 to two times higher compared to Γ_{off} for quasi-static failure. This higher value for Γ_{off} results from a wider damage zone and a higher overall fracture density. The first order estimate for damage zone width in section 4.3.2 suggests that the damage zone width is controlled by rupture velocity. A similar quantitative estimate for the cause of the difference in damage intensity cannot be achieved so easily. First, elastodynamic rupture models predict that with increasing rupture velocity, the state of stress in the rupture tip process zone exceeds the strength of the damage zone rock by an increasing amount [Poliakov et al., 2002; Rice et al., 2005], which may result in the formation of more microfractures. An increasing rupture velocity also increases off-fault strain rates that, when sufficiently high, give rise to a higher microfracture density due to inertia effects [Glenn and Chudnovsky, 1986; Liu et al., 1998; Bhat et al., 2012; Aben et al., 2017]. Second, off-fault stresses arising from slip along a rough fault will cause additional microfracturing and slip along off-fault microfractures formed during rupture. This effect could be represented in the rough fault model by decreasing the off-fault coefficient of friction, which also increases the distance at which slip along a rough fault interacts with rupture-induced off-fault damage. The microstructures of the dynamically failed sample indeed show a few small patches of fine material along secondary fractures at over 10 mm distance from the main fault, indicating that some slip occurred along this secondary fault. Energy dissipation by slip along off-fault microfractures is not considered in the above calculation of Γ_{off} .

Moore and Lockner [1995] observed peak fracture densities of the order of 40–80 mm/mm^2 in the microstructures of a quasi-static rupture propagation experiment on intact Westerly granite at 50 MPa confining pressure. Fracture densities dropped to a background density of around 14 mm/mm^2 at the damage zone boundary defined at 40 mm from the fault. Continuous microstructural observations were limited to 10 mm from the failure zone, except for one measurement at 40 mm distance. Moore and Lockner [1995] report values for Γ_{off} that range from 1.7 to 8.6 kJm^2 , which is similar to the values reported here (2 to 10 kJm^2 , Figure 11c). However, our results show a damage zone width after quasi-static failure of around 10 mm. The similarity in Γ_{off} and the difference in damage zone width is not caused by a difference in resolution; both this study as well as Moore and Lockner [1995] have a cut-off for fractures smaller than 3 μm . Possible explanations for the difference in damage zone widths are: 1) Field and laboratory studies describe the evolution of fracture density in the fault damage zone by an exponential decay [Mitchell and Faulkner, 2009; Faulkner et al., 2011; Moore and Lockner, 1995], a logarithmic decay [Zang et al., 2000], or a powerlaw decay [Savage and Brodsky, 2011; Mayolle et al., 2019; Ostermeijer et al., 2020] with increasing fault-perpendicular distance. This may result in different damage zone widths, but does not affect Γ_{off} much as the ‘tail’ of the damage

zone does not contribute significant amounts of additional fracture damage. 2) This study and *Moore and Lockner* [1995] used different granitic samples. 3) The confining pressure used by *Moore and Lockner* [1995] is half of that used in this study. Earthquake rupture simulations show that the damage zone width decreases with increasing confining pressure, while the relative damage intensity within the damage zone increases [*Okubo et al.*, 2019]. The results from this study and *Moore and Lockner* [1995] comply with these findings: An increase in confining pressure reduces the damage zone width while Γ_{off} remains the same, which equals a higher microfracture intensity in a narrower damage zone.

4.3.4. Off-fault dissipated energy and rupture energetics

Is Γ_{off} a significant energy sink for all preexisting fault in the brittle crust? The primary prerequisite for dissipation of fracture energy in the off-fault volume is that the imposed far-field stresses plus extraneous transient stresses in the rupture tip process zone are sufficiently high to damage the host material. In the experiments presented here, the failure zone material consists of the same material (intact granite) as the host rock, so that strength of the fault interface is the same as the strength of the surrounding material. The imposed stress state during rupture is thus high relative to the strength of the host rock. The magnitude of the additional transient stress field of the rupture tip process zone is proportional to $\Delta\sigma_{ij} \propto \Gamma^{1/2}$ for the limiting case of a singular shear crack [*Freund*, 1990], where Γ is relatively high for intact granite. From these two arguments it follows that stresses around the experimental ruptures are high enough to induce off-fault damage, but should be considered an upper bound for pre-existing fault zones in terms of Γ and strength. We can establish a lower bound scenario for a strong host rock and a weak interface, comprised off two bare granite slabs pressed together. Values of $\Gamma = 0.01 - 3.5 \text{ Jm}^{-2}$ have been published for such an experimental setup [*Ke et al.*, 2018; *Kammer and McLaskey*, 2019]. These values are 5 to 7 orders of magnitude lower than for intact granite and were measured at 6 MPa normal stress, two orders of magnitude lower than our experiment. Thus both imposed far-field stress and the transient stress field are much lower than in our experiment, whereas the off-fault host material remains the same. We therefore expect no off-fault damage and a negligible value for Γ_{off} in these experiments. These two cases mark the extremes for pre-existing faults, were our experiments are more illustrative for faults below 3 km depth where fault core materials likely experience rapid recovery of cohesion by sealing and healing processes, so that the fracture energy Γ of the material increases sufficiently to entice damage in the host rock.

Fracture energy Γ is a material parameter that is independent of fault slip and increases slightly with rupture velocity in most materials (i.e., the change in Γ remains within the same order of magnitude) [*Green and Pratt*, 1974; *Freund*, 1990]. Γ determined from mode I rupture experiments performed in PMMA and glass provide analogue results for mode II shear rupture experiments performed here. During mode I rupture in PMMA and glass, Γ remains more or less constant below a critical velocity that is 0.36 (PMMA) or 0.42 (glass) of the Rayleigh wave speed, but increases by up to a factor 10 at higher rupture velocities up to the Rayleigh wave speed [*Sharon et al.*, 1996]. This increase in Γ is an apparent one caused by microbranching instabilities along the main crack that creates additional fracture surface and accounts for the increase in Γ [*Sharon et al.*, 1996]. At these rupture velocities, the single cracks still obey the initial Γ measured at low rupture velocity [*Sharon and Fineberg*, 1999]. Here, we show that part of the dependence of Γ on rupture velocity is caused by an increasing amount of off-fault dissipated energy Γ_{off} . Γ_{off} itself increases because the off-fault area in which energy is dissipated by microfracturing increases, and the amount of fractures within this area increases as well. What we measure as Γ_{off} in our experiment is qualitatively similar to the additional energy dissipated by microbranching instabilities measured in PMMA during mode I rupture – with the main difference that microbranching around a shear rupture in granite occurs already at quasi-static conditions as evidenced by the off-fault microfractures after quasi-static rupture.

Fracture energy on the main failure plane ($\Gamma - \Gamma_{\text{off}}$) is partly invested as surface energy to create gouge and cataclases, and partly

dissipated as heat. We assume that fracture energy spent on the main failure plane does not change with increasing rupture velocity. Γ thus only increases with rupture velocity if Γ_{off} increases. In our experiments, Γ_{off} doubles from around 3 kJm^{-2} for quasi-static rupture to at least 5.5 kJm^{-2} for dynamic rupture, and so Γ increases by 10%. An increase in Γ means that ruptures will consume more energy to propagate, and a propagating rupture in a material with a velocity-dependent fracture energy will have a decreasing acceleration rate with increasing rupture velocity [*Freund*, 1990].

Although the rupture velocity for the dynamic failure experiment is unknown, we can make a prediction for the evolution of Γ if we adopt the simple relation that Γ_{off} increases linearly with rupture-induced damage zone width. We observe this in our experiments (Figure 11). We then take the relation between rupture velocity and damage zone width (Figure 13), so that we can predict Γ_{off} . Γ_{off} for rupture velocities near the Rayleigh wave speed increases by up to a factor of 10-20 relative to Γ_{off} at low rupture velocity. Near the Rayleigh wave speed, we then expect that $\Gamma = 54 - 80 \text{ kJm}^{-2}$. The factor 10-20 increase in Γ is similar to that measured for PMMA. The critical velocity for a strong increase in Γ for shear failure in granite under confinement is concurrent with the strong increase in damage zone width, at 0.81 of the Rayleigh wave speed ($0.75 V_S$) whereas the critical branching speed for PMMA is 0.36 in mode I rupture.

Field observations show that damage zone width scales linearly with total fault displacement below 1.5-4 km [*Shipton et al.*, 2006; *Savage and Brodsky*, 2011; *Faulkner et al.*, 2011] (Figure 15). These studies argue that this relation is mainly due to slip-related off-fault damage by fault zone roughness and secondary faulting. By approximating off-fault stresses during rupture and during slip along rough faults, we show that for small displacements the rupture tip process zone determines the damage zone width (Figure 13). Our observed damage zone widths after quasi-static (around 10 mm wide) and dynamic rupture (around 20 mm wide) confirm this: They are an order of magnitude larger than the slip that was accumulated during failure (0.83 mm slip for quasi-static rupture and around 3 mm slip for dynamic rupture), and thus do not fit with the linear scaling relation between damage zone width and slip (Figure 15).

This is in contrast with what is argued by *Faulkner et al.* [2011], where it was suggested that the scaling relation goes through the

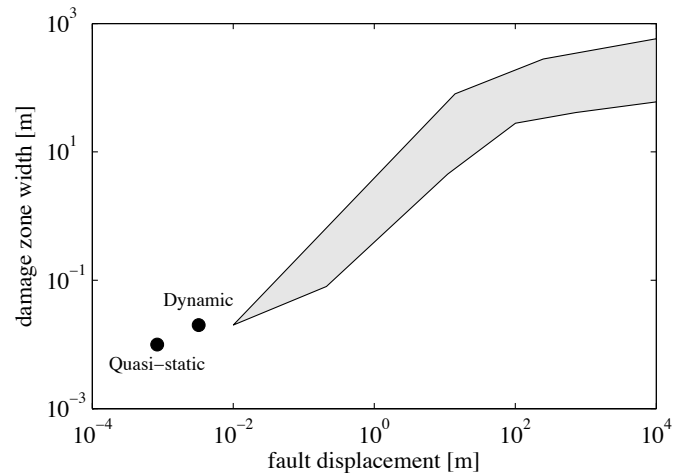


Figure 15: Damage zone width versus total fault displacement, showing the damage zone width and slip from quasi-static rupture LN5 and dynamic rupture experiment LN7 (black datapoints). The shaded area shows the linear scaling relation between damage zone width and displacement, based on field data from *Savage and Brodsky* [2011]; *Faulkner et al.* [2011].

origin (i.e., a zero displacement shear crack has no damage zone). Even at smaller negligible displacements, such as the failed secondary rupture in the mixed rupture experiment, a damage zone width is visible in the P -wave velocity structure (Figure 6). We propose that the lower bound for the scaling relation observed in the field is determined by the stress field around a propagating rupture tip. The absolute value of this lower bound depends on the material properties, far-field stresses, and most importantly the rupture velocity.

Aben *et al.* [2019] argued that the ratio between breakdown work W_b and its off-fault dissipated energy component is proportional to $\delta^{1-\lambda}$. $\lambda \approx 2$ for small earthquake slip below 10 cm, and $\lambda < 1$ for larger slip [Viesca and Garagash, 2015], so that this ratio initially decreases with earthquake slip, to then stabilise or slightly increase with earthquake slip. For quasi-static failure and small slip (< 1 mm for our quasi-static rupture experiment), all the breakdown work is spent as fracture energy, and so $\Gamma_{\text{off}}/\Gamma = 0.1$ [Aben *et al.*, 2019]. However, the scaling proposed by Aben *et al.* [2019] is based on the assumption that damage zone width increases linearly with fault slip as seen in the field [Faulkner *et al.*, 2011; Savage and Brodsky, 2011], whereas our results suggest that at very small amounts of slip the damage zone width is determined by rupture velocity (Figure 15 and 13). The scaling relation between breakdown work and total off-fault dissipated energy is thus only valid when the damage zone width is determined by slip, i.e., fault roughness.

5. Conclusions

We performed dynamic, quasi-static, and mixed shear failure experiments on Lanhélin granite to quantify the off-fault damage in the rupture tip process zones. The in situ P -wave structure and evolution was revealed by laboratory-scale seismic tomography during and after quasi-static failure and after dynamic failure. In both quasi-static and dynamic cases a localised low velocity zone formed around the fault interface, where a maximum reduction in P -wave velocity of about 25% was observed. The low velocity zone around a fault created by dynamic rupture has a similar drop in P -wave velocities. The low velocity zones are caused by off-fault microfractures within the host rock around the fault during rupture and slip. Using an effective medium approach, we computed microfracture densities from the P -wave tomography across the quasi-static and dynamic failure zones. The resulting theoretical microfracture densities are in good agreement with microfracture densities measured from thin sections, indicating that the P -wave tomography reveals realistic near-fault changes in elastic properties. We propose that a similar exercise using high resolution geophysical measurements combined with microstructural measurements from the field can reveal the physical properties around larger fault zones. The damage zone width established from microstructural analysis corresponds to the width of the low P -wave velocity zones, and is around 1 cm wide in the quasi-statically failed sample and 2 cm in the dynamically failed sample. Comparison with a previous microstructural study on quasi-static failed samples suggests that the damage zone width is depth dependent. We argue that the damage zone width in our experiment is controlled by rupture velocity and not by the slip up to a few mm. We propose that at larger slip the damage zone width is determined by fault roughness. Hence, in our experiments the increase in off-fault dissipated energy is mostly caused by an increase in rupture velocity. The off-fault dissipated energy Γ_{off} that we measure is therefore associated to the fracture energy Γ , and was calculated from microstructural observations. Γ_{off} increases from around 3 kJm^{-2} for quasi-static rupture to at least 5.5 kJm^{-2} for dynamic rupture, and shows that shear fracture energy in crystalline material increases with increasing rupture velocity.

Acknowledgments. This study was funded by the UK Natural Environmental Research Council, grants NE/K009656/1 to N.B. and NE/M004716/1 to T.M.M. and N.B., and the European Research Council under the European Union's Horizon 2020 research and innovation

programme (project RockDEaF, grant agreement #804685 to N.B.). We thank J. Davy for thin section preparation and assistance with the SEM. We thank D. Kammer and an anonymous reviewer for their insightful comments. All data needed to evaluate the conclusions in the paper can be found at the NGDC repository of the British Geological Survey (<https://www.bgs.ac.uk/services/NGDC>) in dataset ID128186 for ultrasonic and mechanical data of sample LN5, and dataset ID135445 for all other data and SEM images.

References

- Aben, F. M., M.-L. Doan, J.-P. Gratier, and F. Renard, Coseismic damage generation and pulverization in fault zones: insights from dynamic Split-Hopkinson pressure bar experiments, in *Fault Zone Dynamic Processes: Evolution of Fault Properties During Seismic Rupture*, *Geophys. Monogr. Ser.*, vol. 227, edited by M. Y. Thomas, T. M. Mitchell, and H. S. Bhat, pp. 47–80, American Geophysical Union, Washington, DC, 2017.
- Aben, F. M., N. Brantut, T. M. Mitchell, and E. C. David, Rupture energetics in crustal rock from laboratory-scale seismic tomography, *Geophys. Res. Lett.*, 46, 7337–7344, doi:10.1029/2019GL083040, 2019.
- Aben, F. M., M.-L. Doan, and T. M. Mitchell, Variation of hydraulic properties due to dynamic fracture damage: Implications for fault zones, *J. Geophys. Res.*, 125, e2019JB018,919, 2020.
- Abercrombie, R. E., and J. R. Rice, Can observations of earthquake scaling constrain slip weakening?, *Geophys. J. Int.*, 162, 406–424, 2005.
- Andrews, D. J., Rupture propagation with finite stress in antiplane strain, *J. Geophys. Res.*, 81(20), 3575–3582, 1976.
- Andrews, D. J., Rupture dynamics with energy loss outside the slip zone, *J. Geophys. Res.*, 110, B01307, doi:10.1029/2004JB003191, 2005.
- Atkinson, B. K., Introduction to fracture mechanics and its geophysical applications, in *Fracture mechanics of rock*, edited by B. K. Atkinson, pp. 1–26, Academic Press, London, 1987.
- Barras, F., M. Aldam, T. Roch, E. A. Brener, E. Bouchbinder, and J.-F. Molinari, The emergence of crack-like behavior of frictional rupture: Edge singularity and energy balance, *Earth Planet. Sci. Lett.*, 531, 115,978, 2020.
- Beeler, N., Stress drop with constant, scale independent seismic efficiency and overshoot, *Geophys. Res. Lett.*, 28(17), 3353–3356, 2001.
- Ben-Zion, Y., and J.-P. Ampuero, Seismic radiation from regions sustaining material damage, *Geophys. J. Int.*, 178(3), 1351–1356, 2009.
- Ben-Zion, Y., F. L. Vernon, Y. Ozakin, D. Zigone, Z. E. Ross, H. Meng, M. White, J. Reyes, D. Hollis, and M. Barklage, Basic data features and results from a spatially dense seismic array on the San Jacinto fault zone, *Geophys. J. Int.*, 202(1), 370–380, 2015.
- Bhat, H. S., A. J. Rosakis, and C. G. Sammis, A micromechanics based constitutive model for brittle failure at high strain rates, *J. Appl. Mech.*, 79, 031,016.1–031,016.12, 2012.
- Brantut, N., Time-dependent recovery of microcrack damage and seismic wave speeds in deformed limestone, *J. Geophys. Res.*, 120(12), 8088–8109, doi:10.1002/2015JB012324, 2015.
- Brantut, N., Time-resolved tomography using acoustic emissions in the laboratory, and application to sandstone compaction, *Geophys. J. Int.*, 213, 2177–2192, 2018.
- Brantut, N., Dilatancy-induced fluid pressure drop during dynamic rupture: Direct experimental evidence and consequences for earthquake dynamics, *Earth Planet. Sci. Lett.*, 538, 116,179, 2020.
- Brantut, N., and T. M. Mitchell, Assessing the efficiency of thermal pressurisation using natural pseudotachylite-bearing rocks, *Geophys. Res. Lett.*, 45, 9533–9541, doi:10.1029/2018GL078649, 2018.
- Brantut, N., and R. C. Viesca, The fracture energy of ruptures driven by flash heating, *Geophys. Res. Lett.*, 44, 6718–6725, doi:10.1002/2017GL074110, 2017.
- Brantut, N., A. Schubnel, and Y. Guéguen, Damage and rupture dynamics at the brittle-ductile transition: the case of gypsum, *J. Geophys. Res.*, 116, B01404, doi:10.1029/2010JB007675, 2011.
- Brantut, N., M. J. Heap, P. Baud, and P. G. Meredith, Rate- and strain-dependent brittle deformation of rocks, *J. Geophys. Res.*, 119, doi: 10.1002/2013JB010448, 2014.
- Cappa, F., C. Perrin, I. Manighetti, and E. Delor, Off-fault long-term damage: A condition to account for generic, triangular earthquake slip, *Geochem. Geophys. Geosyst.*, 15, 1476–1493, 2014.
- Chester, F. M., and J. S. Chester, Stress and deformation along wavy frictional faults, *J. Geophys. Res.*, 105(B10), 23,421–23,430, 2000.
- Chester, J. S., F. M. Chester, and A. K. Kronenberg, Fracture surface energy of the Punchbowl fault, San Andreas system, *Nature*, 437(1), 133–136, 2005.

- Dieterich, J. H., and D. E. Smith, Non-planar faults: Mechanics of slip and off-fault damage, *Pure Appl. Geophys.*, 166, 1799–1815, 2009.
- Dunham, E. M., D. Belanger, L. Cong, and J. E. Kozdon, Earthquake ruptures with strongly rate-weakening friction and off-fault plasticity, Part 1: Planar faults, *Bull. Seism. Soc. Am.*, 101(5), 2296–2307, 2011.
- Eccles, D., P. R. Sammonds, and O. C. Clint, Laboratory studies of electrical potential during rock fracture, *Int. J. Rock Mech. Min. Sci.*, 42(7–8), 933–949, 2005.
- Faulkner, D. R., T. M. Mitchell, D. Healy, and M. J. Heap, Slip on 'weak' faults by the rotation of regional stress in the fracture damage zone, *Nature*, 444, 922–925, doi:10.1038/nature05353, 2006.
- Faulkner, D. R., T. M. Mitchell, E. Jensen, and J. Cembrano, Scaling of fault damage zones with displacement and the implications for fault growth processes, *J. Geophys. Res.*, 116, B05403, 2011.
- Freund, L. B., *Dynamic Fracture Mechanics*, Cambridge University Press, Cambridge, UK, 1990.
- Gavrilenko, P., and Y. Guéguen, Pressure dependence of permeability: a model for cracked rocks, *Geophys. J. Int.*, 98, 159–172, 1989.
- Glenn, L. A., and A. Chudnovsky, Strain-energy effects on dynamic fragmentation, *J. Appl. Phys.*, 59(4), 1379–1380, 1986.
- Green, A. K., and P. L. Pratt, Measurement of the dynamic fracture toughness of polymethylmethacrylate by high-speed photography, *Eng. Frac. Mech.*, 6, 71–80, 1974.
- Griffiths, L., M. J. Heap, P. Baud, and J. Schmittbuhl, Quantification of microcrack characteristics and implications for stiffness and strength of granite, *Int. J. Rock Mech. Min. Sci.*, 100, 138–150, 2017.
- Guéguen, Y., and M. Kachanov, Effective elastic properties of cracked and porous rocks – an overview –, in *Mechanics of Crustal Rocks*, edited by Y. M. Leroy and F. K. Lehner, pp. 73–125, Springer, Berlin Heidelberg, 2011.
- Guéguen, Y., and A. Schubnel, Elastic wave velocities and permeability of cracked rocks, *Tectonophysics*, 370(1–4), 163–176, 2003.
- Hillers, G., P. Roux, M. Campillo, and Y. Ben-Zion, Focal spot imaging based on zero lag cross correlation amplitude fields : Application to dense array data at the San Jacinto fault zone., *J. Geophys. Res.*, 121, 8048–8067, 2016.
- Jeppson, T. N., K. K. Bradbury, and J. P. Evans, Geophysical properties within the San Andreas fault zone at the San Andreas Fault Observatory at Depth and their relationships to rock properties and fault zone structure, *J. Geophys. Res.*, 115, 1–20, 2010.
- Kammer, D. S., and G. C. McLaskey, Fracture energy estimates from large-scale laboratory earthquakes, *Earth Planet. Sci. Lett.*, 511, 36–43, 2019.
- Kanamori, H., and L. Rivera, Energy partitioning during and earthquake, in *Earthquakes: radiated energy and the physics of faulting*, edited by R. Abercrombie, A. McGarr, G. Di Toro, and H. Kanamori, American Geophysical Union, Washington, DC, 2006.
- Ke, C.-Y., G. C. McLaskey, and D. S. Kammer, Rupture termination in laboratory-generated earthquakes, *Geophys. Res. Lett.*, 45, 12,784–12,792, 2018.
- Liu, C., W. G. Knauss, and A. J. Rosakis, Loading rates and the dynamic initiation toughness in brittle solids, *Int. J. Frac.*, 90, 103–118, 1998.
- Lockner, D. A., J. D. Byerlee, V. Kuksenko, A. Ponomarev, and A. Sidorin, Quasi-static fault growth and shear fracture energy in granite, *Nature*, 350(6313), 39–42, 2001.
- Martin, R. J., Pore pressure stabilization of failure in Westerly granite, *Geophys. Res. Lett.*, 7(5), 404–406, 1980.
- Mayolle, S., R. Soliva, Y. Caniven, C. Wibberley, G. Ballas, G. Milesi, and S. Dominguez, Scaling of fault damage zones in carbonate rocks, *J. Struct. Geol.*, 124, 35–50, 2019.
- Mitchell, T. M., and D. R. Faulkner, The nature and origin of off-fault damage surrounding strike-slip fault zones with a wide range of displacements: A field study from the Atacama fault system, northern Chile, *J. Struct. Geol.*, 31, 802–816, 2009.
- Mooney, W. D., and A. Ginzburg, Seismic measurements of the internal properties of fault zones, *Pure Appl. Geophys.*, 124(1/2), 141–157, 1986.
- Moore, D. E., and D. A. Lockner, The role of microcracking in shear-fracture propagation in granite, *J. Struct. Geol.*, 17(1), 95–111, 1995.
- Oda, M., A method for evaluating the effect of crack geometry on the mechanical behavior of cracked rock masses, *Mech. Mat.*, 2, 163–171, 1983.
- Okubo, K., H. S. B. E. Rougier, S. Marty, A. Schubnel, Z. Lei, E. E. Knight, and Y. Klinger, Dynamics, radiation, and overall energy budget of earthquake rupture with coseismic off-fault damage, *J. Geophys. Res.*, 124, 11,771–11,801, 2019.
- Ostermeijer, G. A., T. M. Mitchell, F. M. Aben, M. T. Dorsey, J. Browning, T. K. Rockwell, J. M. Fletcher, and F. Ostermeijer, Damage zone heterogeneity on seismogenic faults in crystalline rock: a field study of the Borrego fault, Baja California, *J. Struct. Geol.*, 137, 104016, 2020.
- Passelègue, F. X., L. X. Pimienta, D. R. Faulkner, A. Schubnel, J. Fortin, and Y. Guéguen, Development and recovery of stress-induced elastic anisotropy during cyclic loading experiment, *Geophys. Res. Lett.*, 45, 8156–8166, 2018.
- Paterson, M. S., and T. F. Wong, *Experimental Rock Deformation – The Brittle Field*, 2nd ed., Springer-Verlag, Berlin Heidelberg, 2005.
- Poliakov, A. N. B., R. Dmowska, and J. R. Rice, Dynamic shear rupture interactions with fault bends and off-axis secondary faulting, *J. Geophys. Res.*, 107(B11), 2295, doi:10.1029/2001JB000572, 2002.
- Qiu, H., Y. Ben-Zion, Z. E. Ross, P.-E. Share, and F. L. Vernon, Internal structure of the San Jacinto fault zone at Jackass Flat from data recorded by a dense linear array, *Geophys. J. Int.*, 3(1), 1369–1388, 2017.
- Reches, Z., and D. A. Lockner, Nucleation and growth of faults in brittle rocks, *J. Geophys. Res.*, 99(94), 18,159–18,173, 1994.
- Rempe, M., T. M. Mitchell, J. Renner, S. Nippres, Y. Ben-Zion, and T. Rockwell, Damage and seismic velocity structure of pulverized rocks near the San Andreas fault, *J. Geophys. Res.*, 118, 2813–2831, doi: 10.1002/jgrb.50184, 2013.
- Rempe, M., T. M. Mitchell, J. Renner, S. A. F. Smith, A. Bistacchi, and G. Di Toro, The relationship between microfracture damage and the physical properties of fault-related rocks: The Gole Larghe fault zone, Italian Southern Alps, *J. Geophys. Res.*, 123, 7661–7687, 2018.
- Rice, J. R., The mechanics of earthquake rupture, in *Physics of the Earth's Interior*, edited by A. M. Dziewonski and E. Boschi, Proc. Intl. School of Physics E. Fermi, pp. 555–649, Italian Physical Society/North Holland Publ. Co., 1980.
- Rice, J. R., and J. W. Rudnicki, A note on some features of the theory of localisation of deformation, *Int. J. Solids Struct.*, 16(7), 597–605, 1980.
- Rice, J. R., C. G. Sammis, and R. Parsons, Off-fault secondary failure induced by a dynamic slip pulse, *Bull. Seism. Soc. Am.*, 95(1), 109–134, 2005.
- Rizzo, R. E., D. Healy, and L. De Siena, Benefits of maximum likelihood estimators for fracture attribute analysis: Implications for permeability and up-scaling, *J. Struct. Geol.*, 95, 17–31, 2017.
- Rockwell, T., M. Sisk, G. Girty, O. Dor, N. Wechsler, and Y. Ben-Zion, Chemical and physical characteristics of pulverized Tejon Lookout granite adjacent to the San Andreas and Garlock faults: Implications for earthquake physics, *Pure Appl. Geophys.*, 166, 1725–1746, 2009.
- Savage, H. M., and E. E. Brodsky, Collateral damage: Evolution with displacement of fracture distribution and secondary fault strands in fault damage zones, *J. Geophys. Res.*, 116(3), B03405, 2011.
- Sayers, C., and M. Kachanov, Microcrack-induced elastic wave anisotropy of brittle rocks, *J. Geophys. Res.*, 100(B3), 4149–4156, 1995.
- Schubnel, A., O. Nishizawa, K. Masuda, X. J. Lei, Z. Xue, and Y. Guéguen, Velocity measurements and crack density determination during wet triaxial experiments on Oshima and Toki granites, *Pure Appl. Geophys.*, 160, 869–887, 2003.
- Segall, P., and J. R. Rice, Dilatancy, compaction, and slip instability of a fluid-infiltrated fault, *J. Geophys. Res.*, 100(B11), 22,155–22,171, 1995.
- Segall, P., A. M. Rubin, A. M. Bradley, and J. R. Rice, Dilatant strengthening as a mechanism for slow slip events, *J. Geophys. Res.*, 115, B12305, doi:10.1029/2010JB007449, 2010.
- Sharon, E., and J. Fineberg, Confirming the continuum theory of dynamic brittle fracture for fast cracks, *Nature*, 397(6717), 333–335, 1999.
- Sharon, E., S. Gross, and J. Fineberg, Energy dissipation in dynamic fracture, *Phys. Rev. Lett.*, 76(12), 2117–2120, 1996.
- Shipton, Z. K., A. M. Soden, J. D. Kirkpatrick, A. M. Bright, and R. J. Lunn, How thick is a fault? Fault displacement-thickness scaling revisited, in *Earthquakes: radiated energy and the physics of faulting*, *Geophys. Monogr. Ser.*, vol. 170, edited by R. Abercrombie, A. McGarr, G. Di Toro, and H. Kanamori, pp. 193–198, American Geophysical Union, Washington, DC, 2006.
- Siratovich, P. A., M. C. Villeneuve, J. W. Cole, B. M. Kennedy, and F. Bégué, Saturated heating and quenching of three crustal rocks and implications for thermal stimulation of permeability in geothermal reservoirs, *Int. J. Rock Mech. Min. Sci.*, 80, 265–280, 2105.
- Stanchits, S. A., D. A. Lockner, and A. V. Ponomarev, Anisotropic changes in *p*-wave velocity and attenuation during deformation and fluid infiltration of granite, *Bull. Seism. Soc. Am.*, 93(4), 1803–1822, 2003.
- Taponnier, P., and W. F. Brace, Development of stress-induced microcracks in Westerly granite, *Int. J. Rock Mech. Min. Sci. & Geomech. Abstr.*, 13, 103–112, 1976.
- Tarantola, A., *Inverse Problem Theory*, 2nd ed., Society for Industrial Mathematics, Philadelphia, 2005.
- Thomas, M. Y., and H. S. Bhat, Dynamic evolution of off-fault medium during an earthquake : A micromechanics based model, *Geophys. J. Int.*, 214(2), 1267–1280, 2018.

- Thomas, M. Y., H. S. Bhat, and Y. Klinger, Effect of brittle off-fault damage on earthquake rupture dynamics, in *Fault Zone Dynamic Processes: Evolution of Fault Properties During Seismic Rupture*, *Geophys. Monogr. Ser.*, vol. 227, edited by M. Y. Thomas, T. M. Mitchell, and H. S. Bhat, pp. 171–194, American Geophysical Union, Washington, DC, 2017.
- Tinti, E., P. Spudich, and M. Cocco, Earthquake fracture energy inferred from kinematic rupture models on extended faults, *J. Geophys. Res.*, *110*, B12303, doi:10.1029/2005JB003644, 2005.
- Viesca, R. C., and D. I. Garagash, Ubiquitous weakening of faults due to thermal pressurization, *Nat. Geosci.*, doi:10.1038/ngeo2554, 2015.
- Wawersik, W. R., and W. F. Brace, Post-failure behavior of a granite and diabase, *Rock Mech.*, *3*, 61–85, 1971.
- Wong, T.-F., Shear fracture energy of Westerly granite from post-failure behavior, *J. Geophys. Res.*, *87*(B2), 990–1000, 1982.
- Wong, T.-F., On the normal stress dependence of the shear fracture energy, in *Earthquake Source Mechanics*, *Geophys. Monogr. Ser.*, vol. 37, edited by S. Das, J. Boatwright, and C. H. Scholz, pp. 1–11, American Geophysical Union, Washington, DC, 1986.
- Xu, S., Y. Ben-Zion, J.-P. Ampuero, and V. Lyakhovskiy, Dynamic ruptures on a frictional interface with off-fault brittle damage: Feedback mechanisms and effects on slip and near-fault motion, *Pure Appl. Geophys.*, *172*(5), 1243–1267, 2015.
- Zang, A., F. C. Wagner, S. Stanchits, C. Janssens, and G. Dresen, Fracture process zone in granite, *J. Geophys. Res.*, *105*(B10), 23,651–23,661, 2000.

F. M. Aben, Department of Earth Sciences, University College London, Gower Street, London WC1E 6BS, UK. (f.aben@ucl.ac.uk)